

Impact cratering as a planetary process: Information from impact glasses

G. Giuli



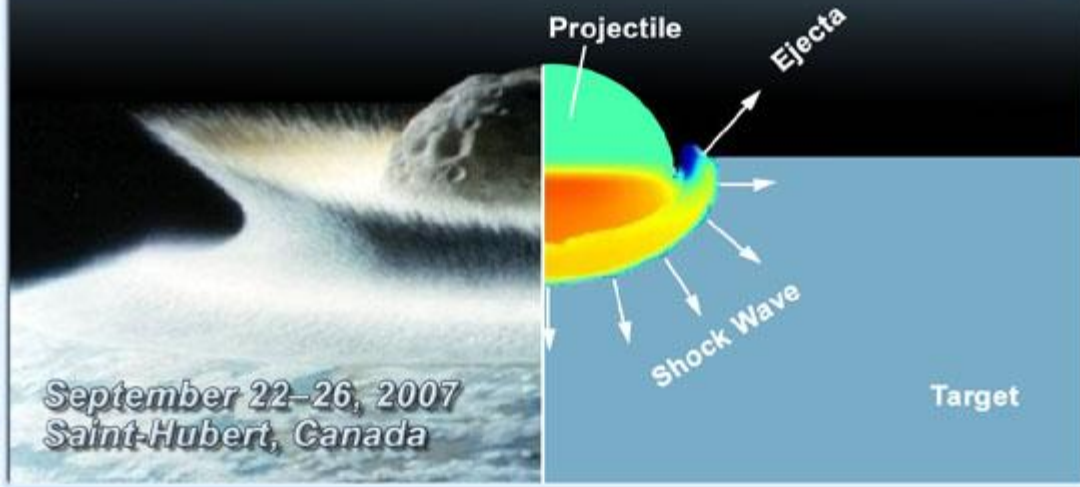
Scuola d



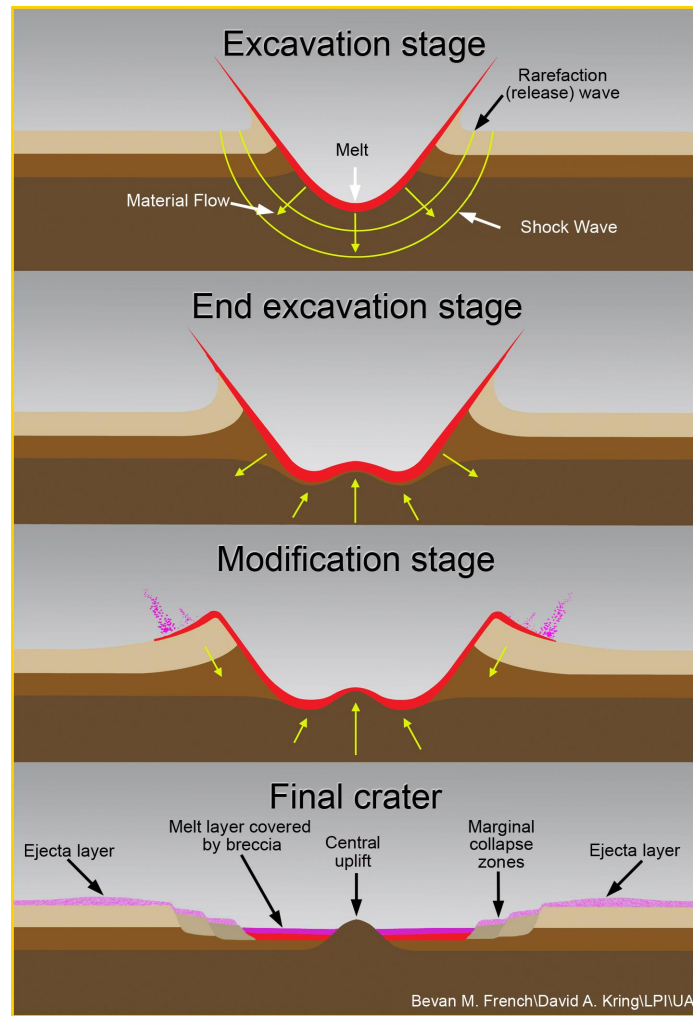
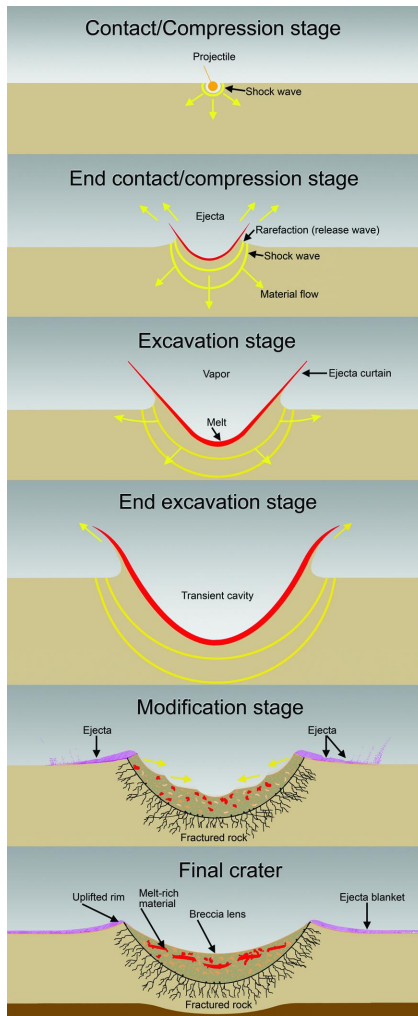
cnologie- sez C... Università di Cam



BRIDGING THE GAP II: EFFECT OF TARGET PROPERTIES ON THE IMPACT CRATERING PROCESS



*September 22–26, 2007
Saint-Hubert, Canada*



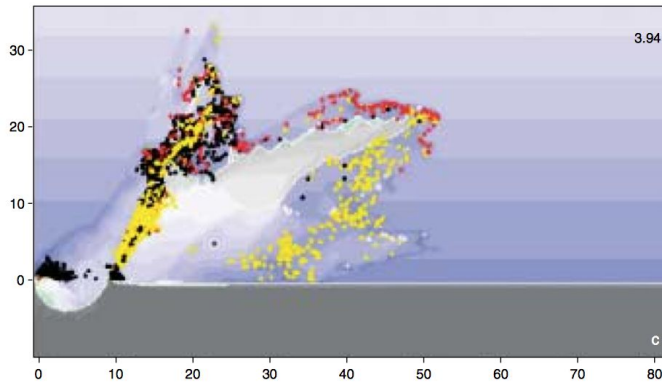
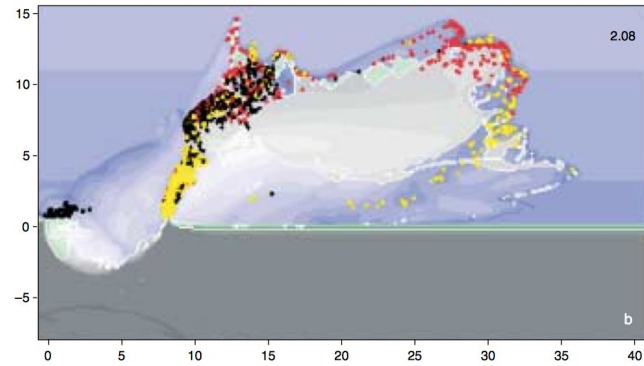
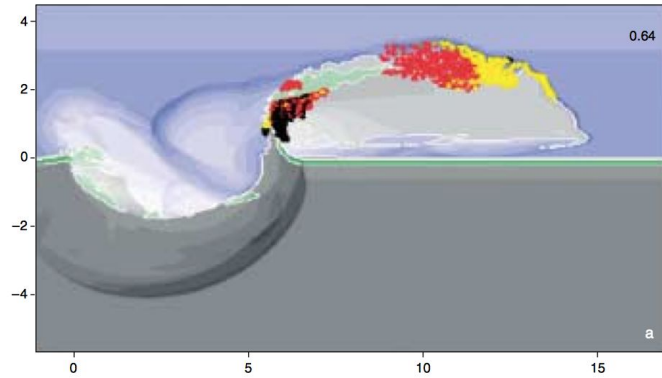
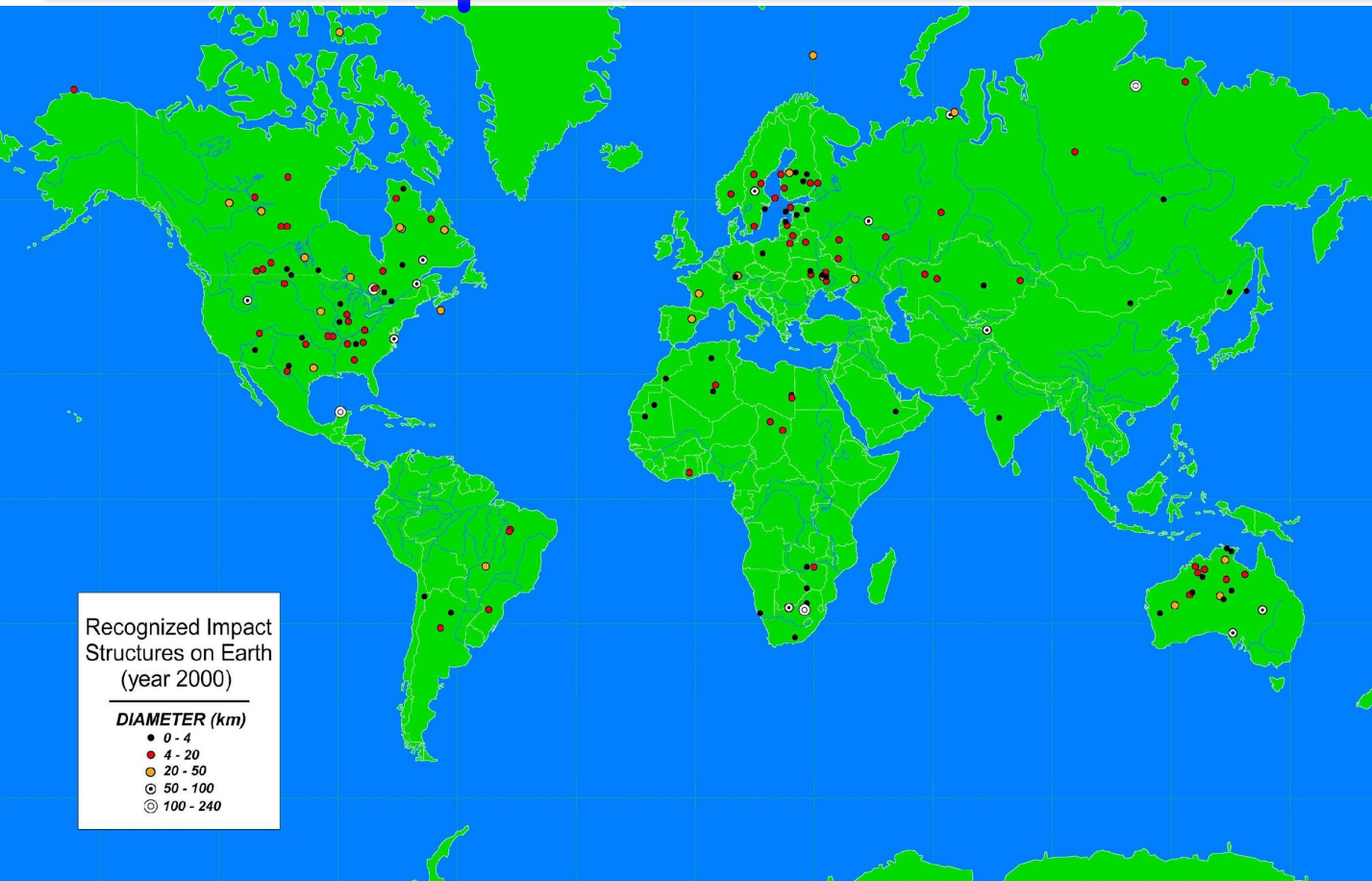


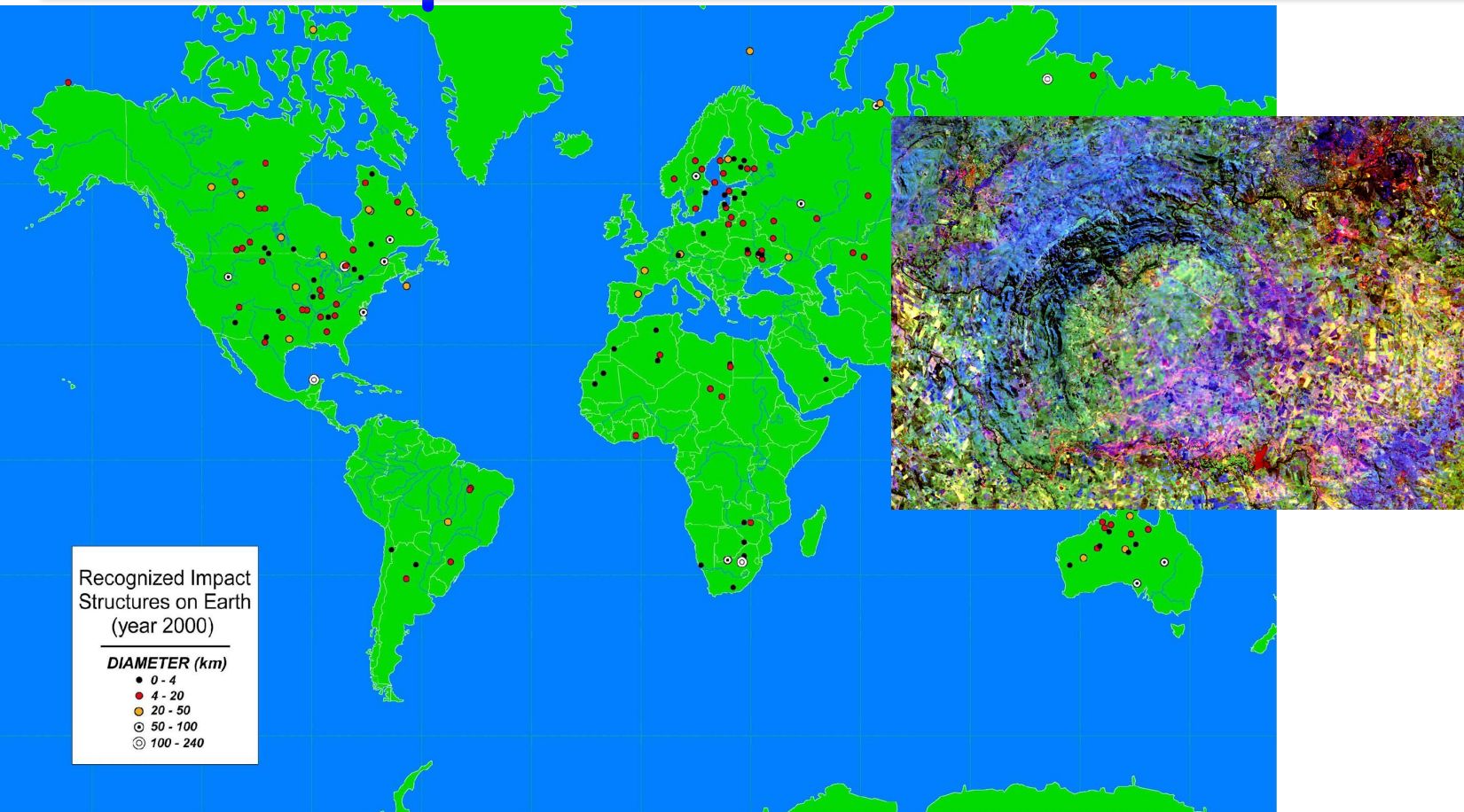
Figure 3. Tektite ejection model results for a 30° impact and impact speed of 20 km/s at a) 0.6 sec, b) 2 sec, and c) 3.9 sec. after the impact. ● – molten upper layer material (possible tektites); ● – molten target materials (not tektites); ● – solid target material.

Ejection model
 Stoffer, Artemieva, Pierazzo (2002) M&PS, 37, 1893-1907

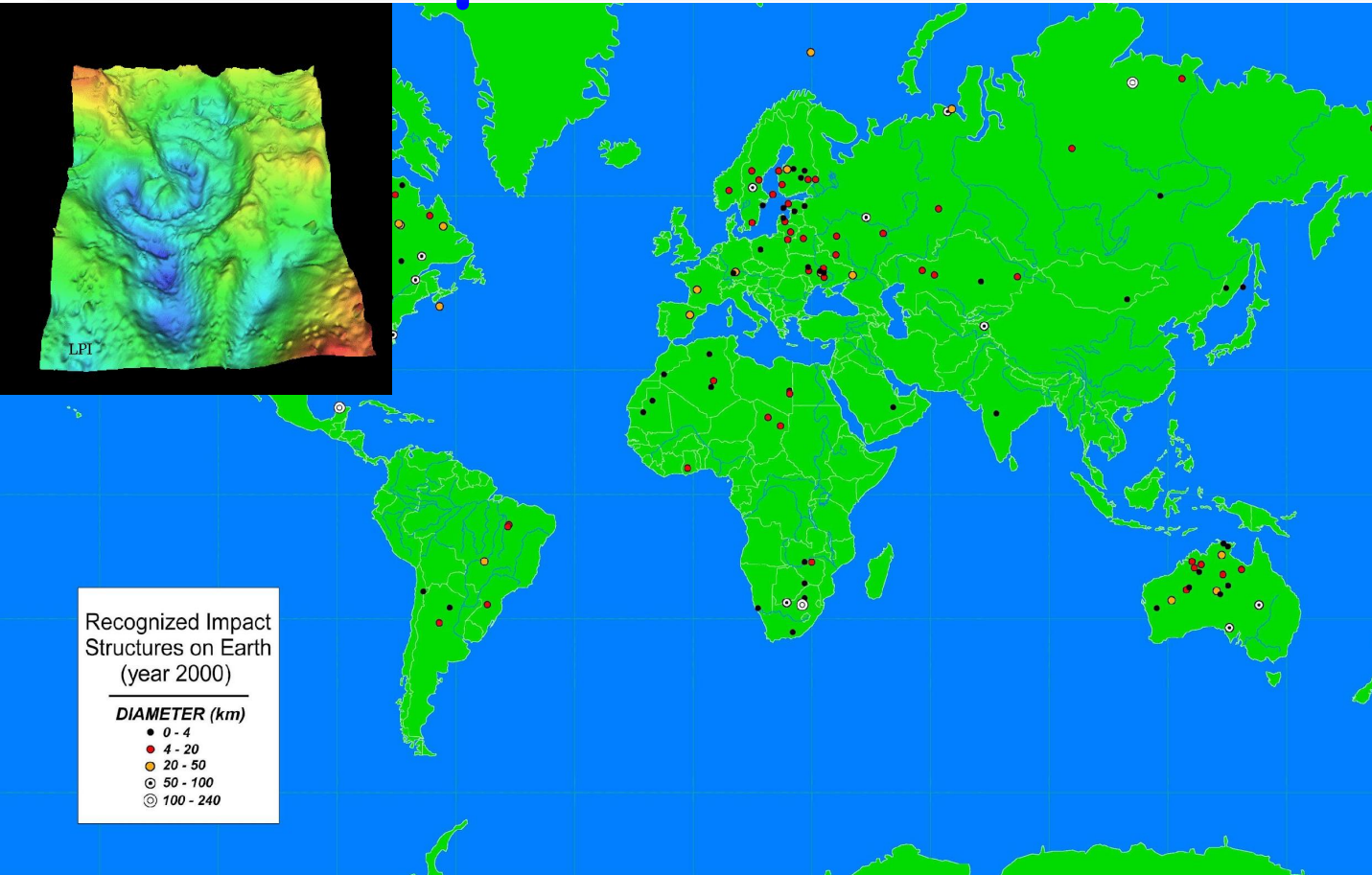
Impact crater distribution



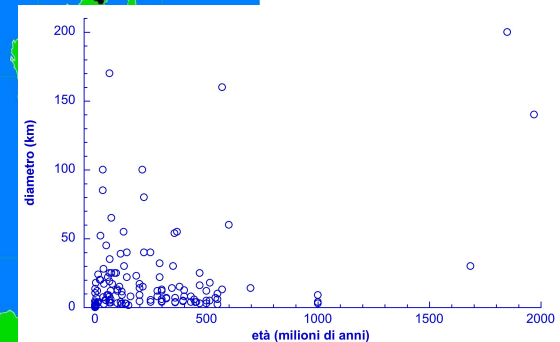
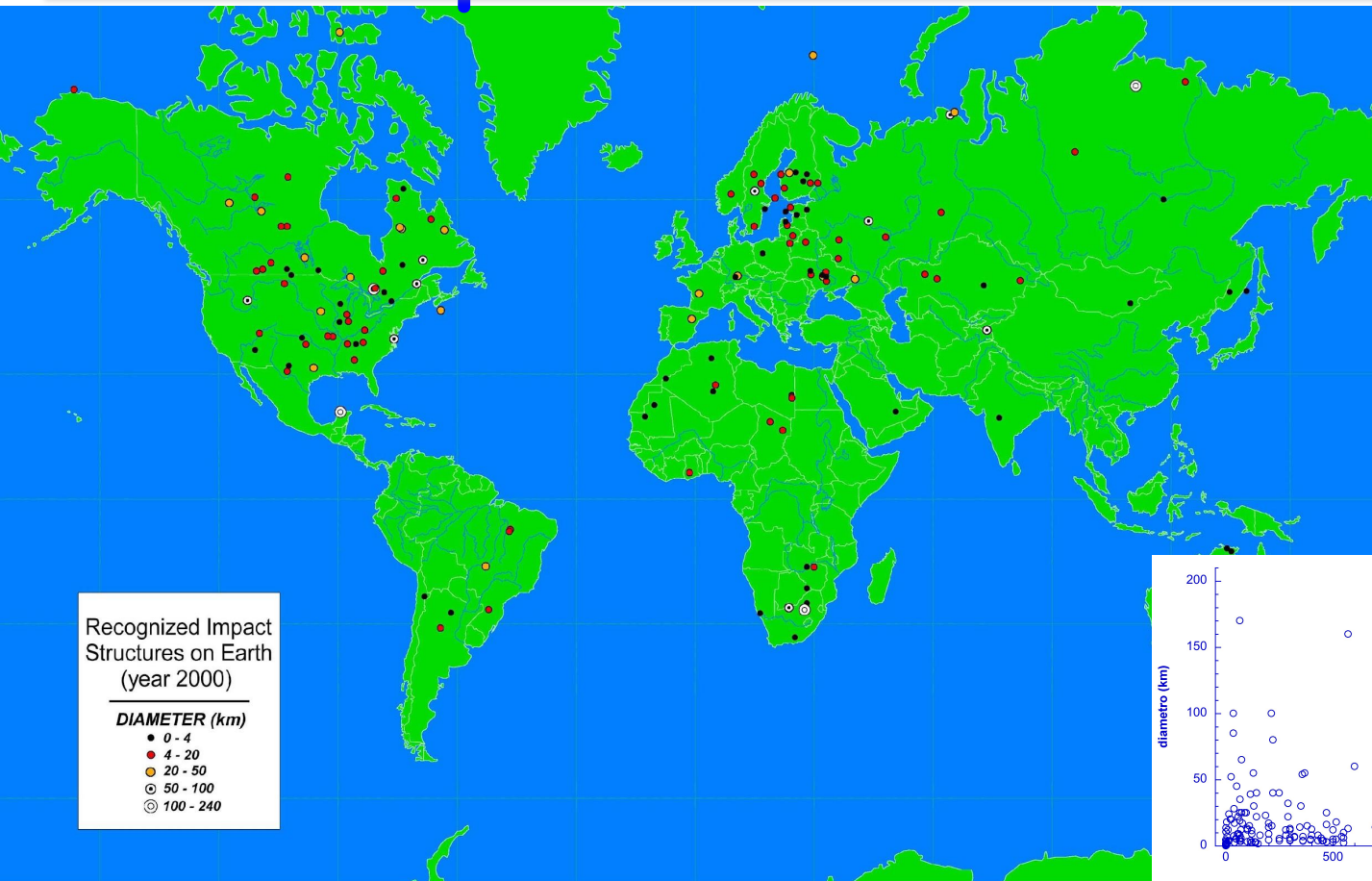
Impact crater distribution

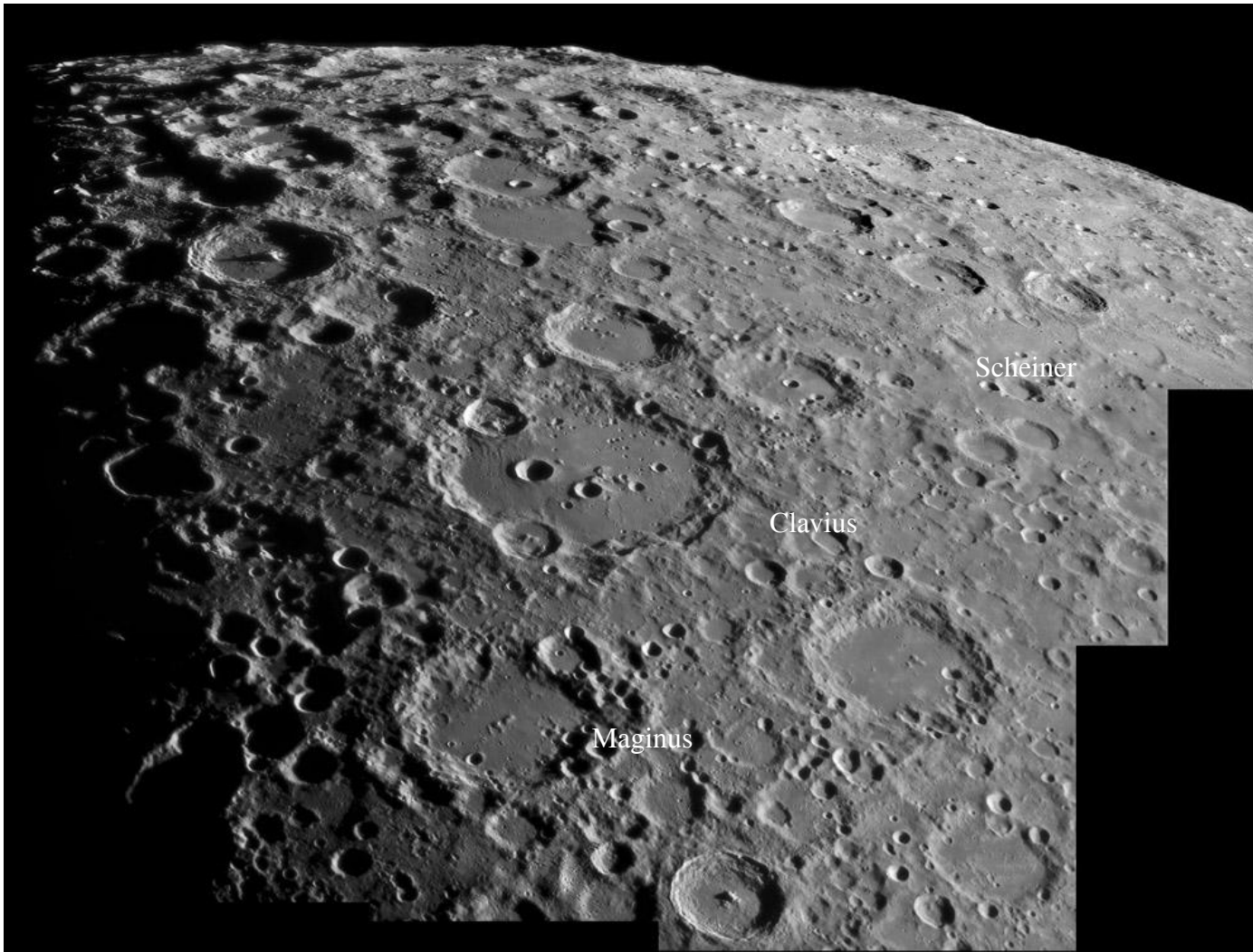


Impact crater distribution



Impact crater distribution

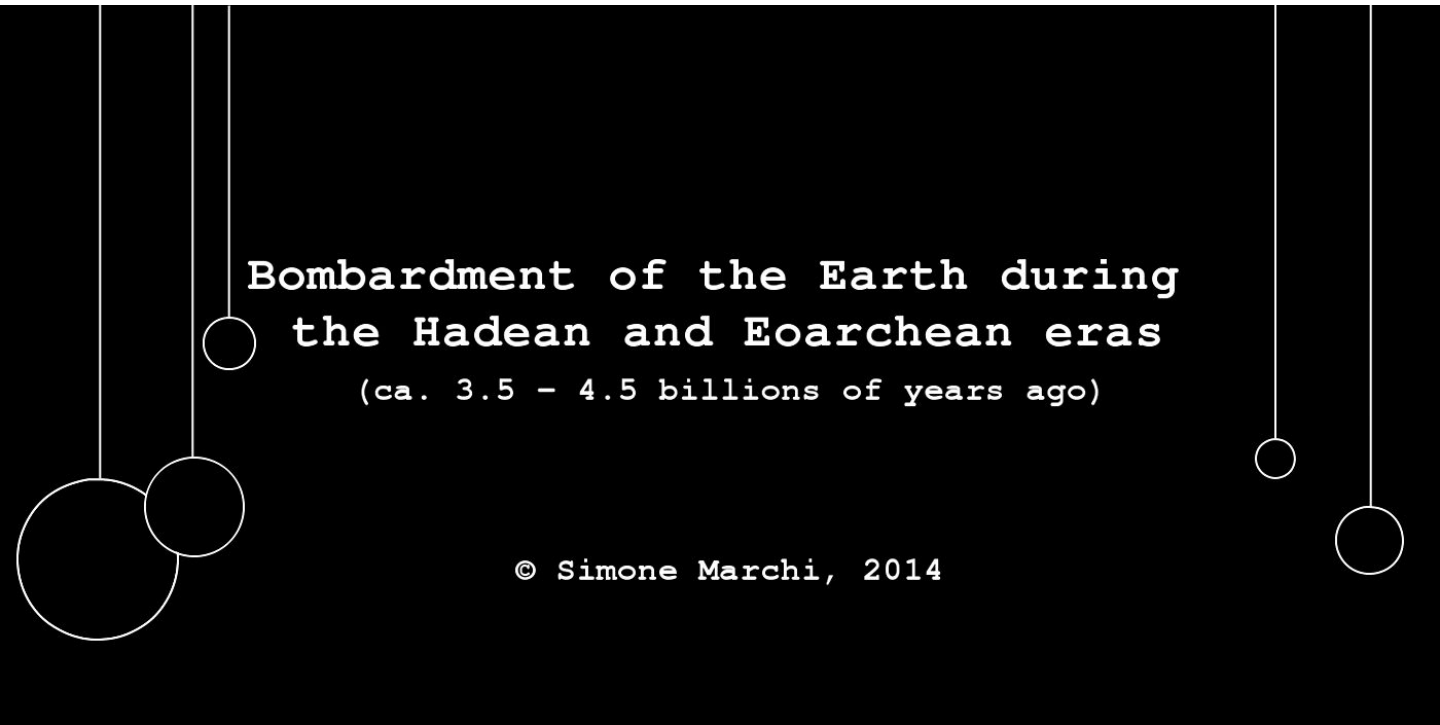




Scheiner

Clavius

Maginus



**Bombardment of the Earth during
the Hadean and Eoarchean eras**
(ca. 3.5 - 4.5 billions of years ago)

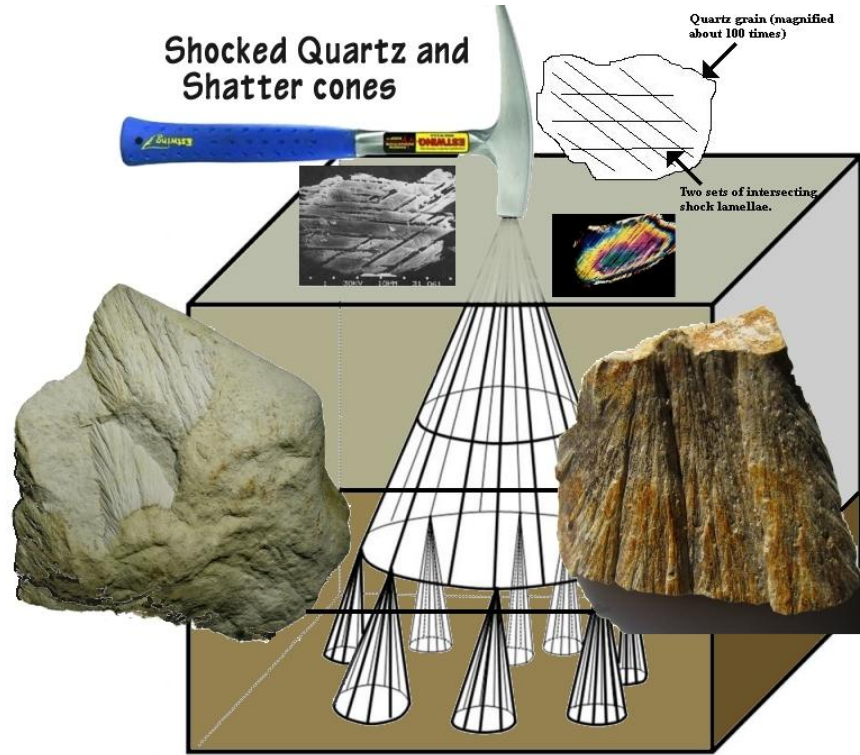
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Megascopic features

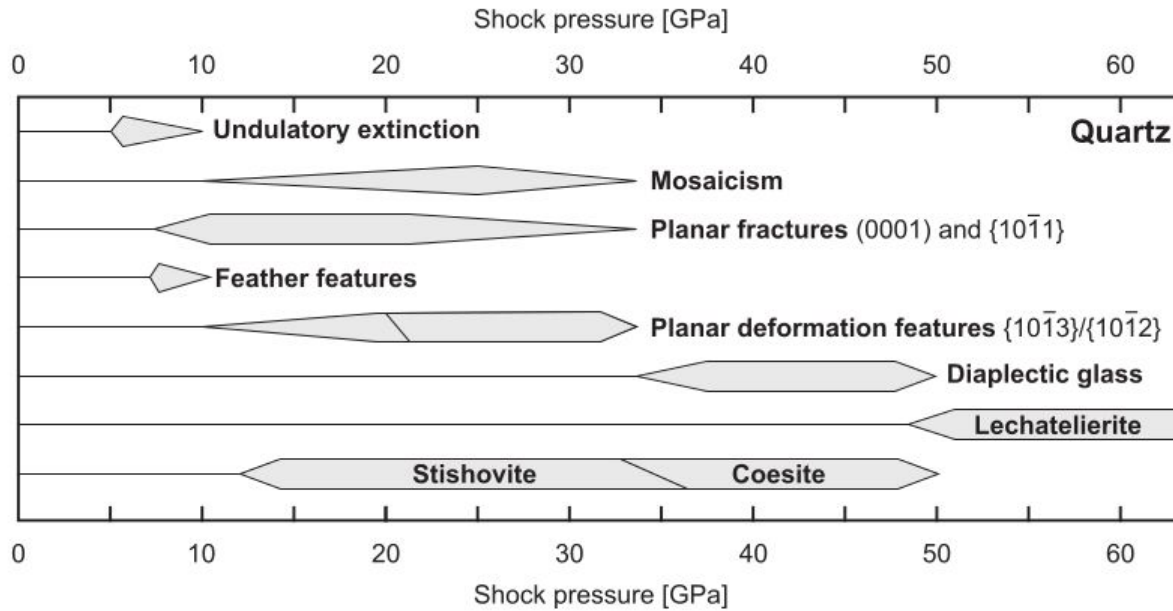


Shatter cones from the Ries crater (D)

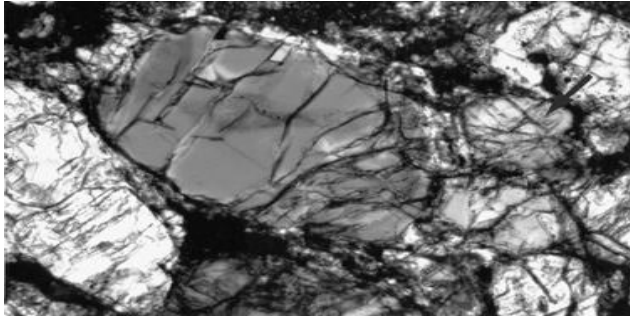
Megascopic features



Microscopic features



-Undulose extinction



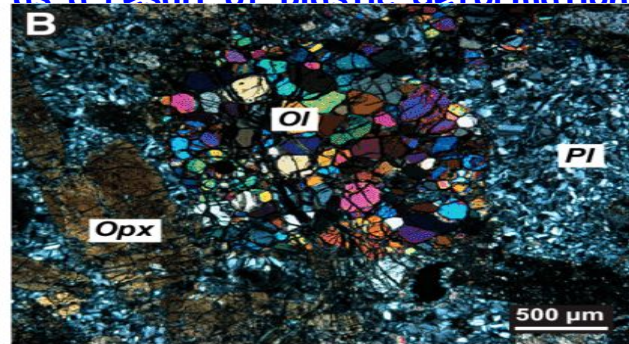
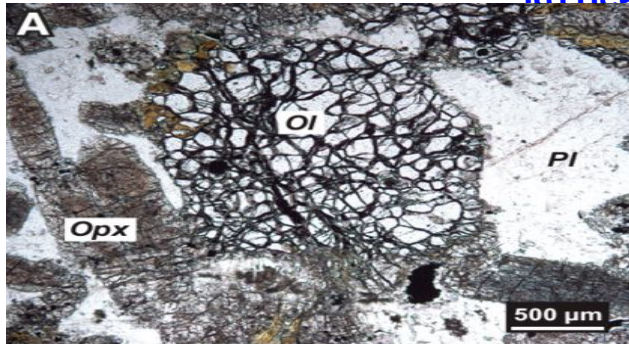
Olivine showing weak undulose extinction
In the framville chondrite (H6).

A crystal showing undulose extinction has somewhat non-uniform extinction caused by slight-to-moderate misorientation of the crystal lattice as a result of Plastic deformation.

-Mosaicism

Olivine showing weak undulose extinction
In the framville chondrite (H6).

A crystal showing mosaic extinction has patchy extinction caused by moderate-to-large misorientation of the crystal lattice as a result of plastic deformation.



-Planar Deformation Features

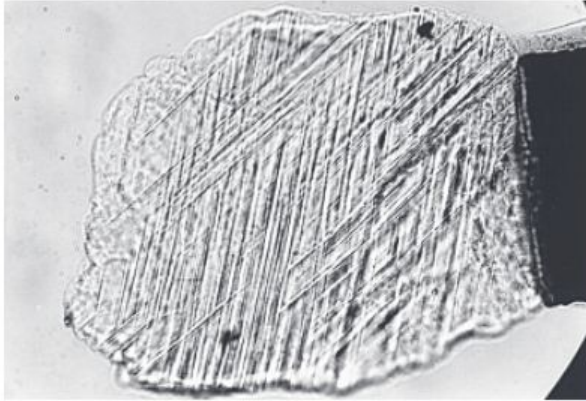


Fig. 4.17. Quartz; multiple PDFs, fresh. Small quartz grain (0.20 mm long) from K/T boundary ejecta layer, showing two prominent sets of fresh (undecorated) PDFs. (Small dots with halos are artifacts.) Specimen from Starkville South, a few kilometers south of Trinidad, Colorado. Photograph courtesy of G. A. Izett. Spindle stage mount (plane-polarized light).

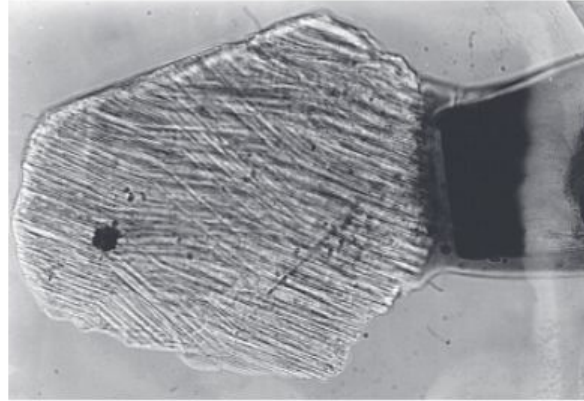


Fig. 4.18. Quartz; multiple PDFs, fresh. Small quartz grain (0.36 mm long) from K/T boundary ejecta layer, containing one opaque inclusion and multiple (3-5?) prominent sets of fresh (undecorated) PDFs. Specimen from Clear Creek North, a few kilometers south of Trinidad, Colorado. Photograph courtesy of G. A. Izett. Spindle stage mount (plane-polarized light).

-Diaplectic glass

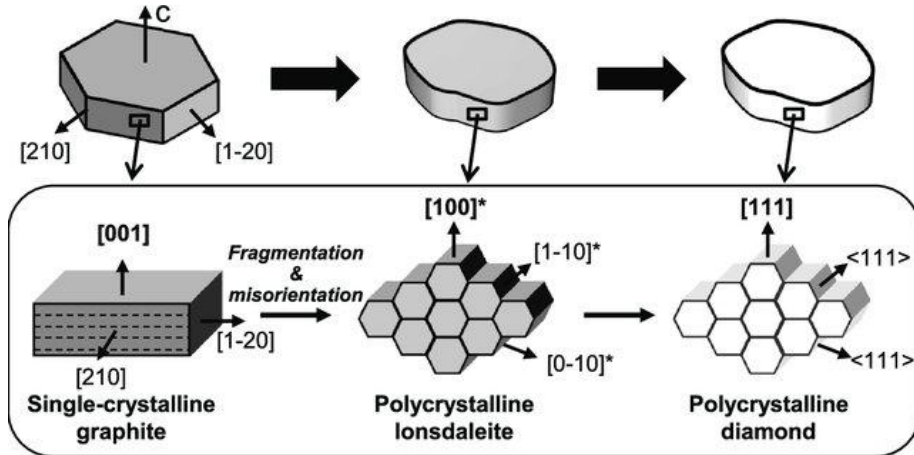
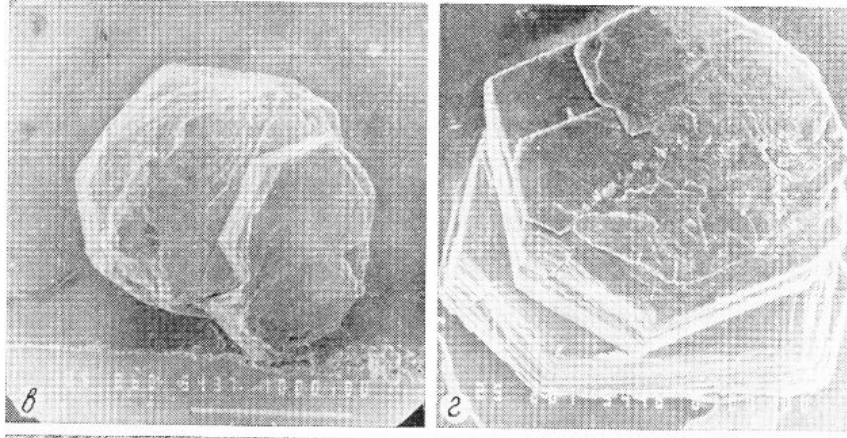


The passage of a shock wave can cause a solid-state amorphization

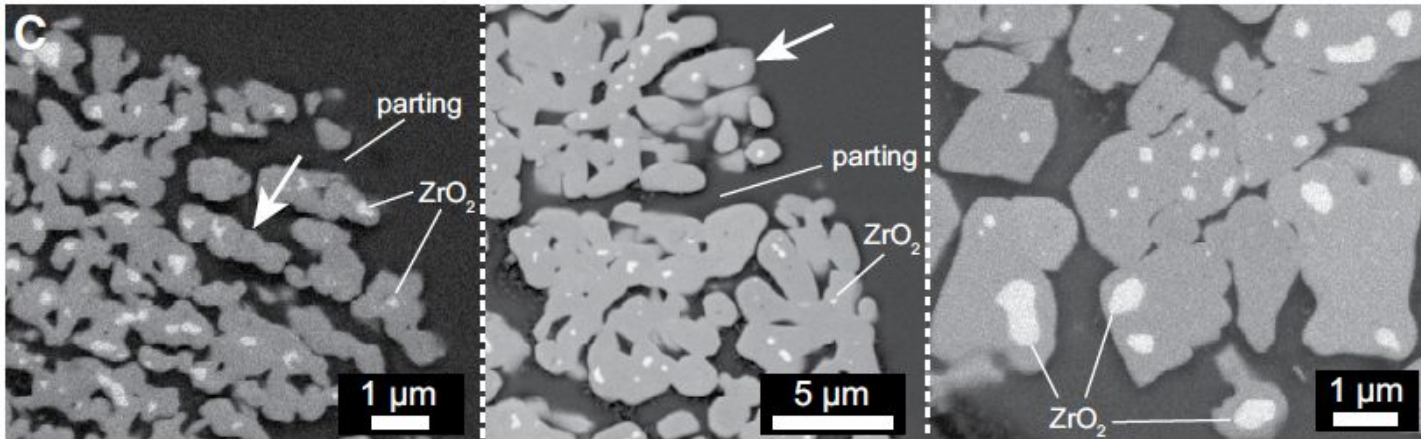
No melting features (flow) observed

Pre-existing fractures and cleavages can be preserved

-Polimorphic transformation Diamond paramorph on graphite



-Thermal decomposition



The heat produced by the passage of a shock wave can trigger a thermal decomposition



e.g. $\text{ZrSiO}_4 \rightarrow \text{ZrO}_2 + \text{SiO}_2$ @ 1673 °C
typical also of fulgurites

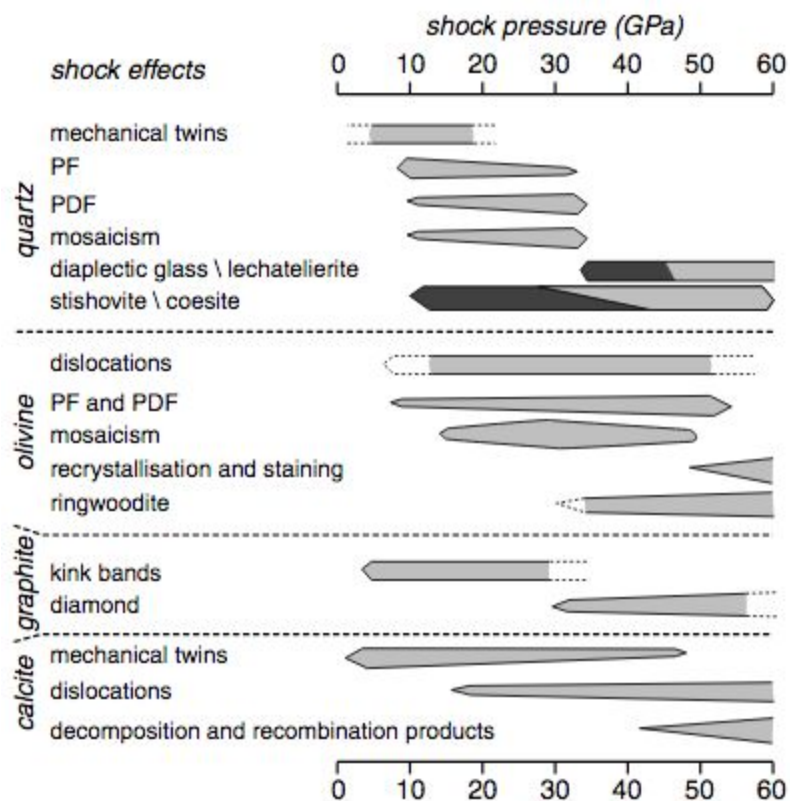
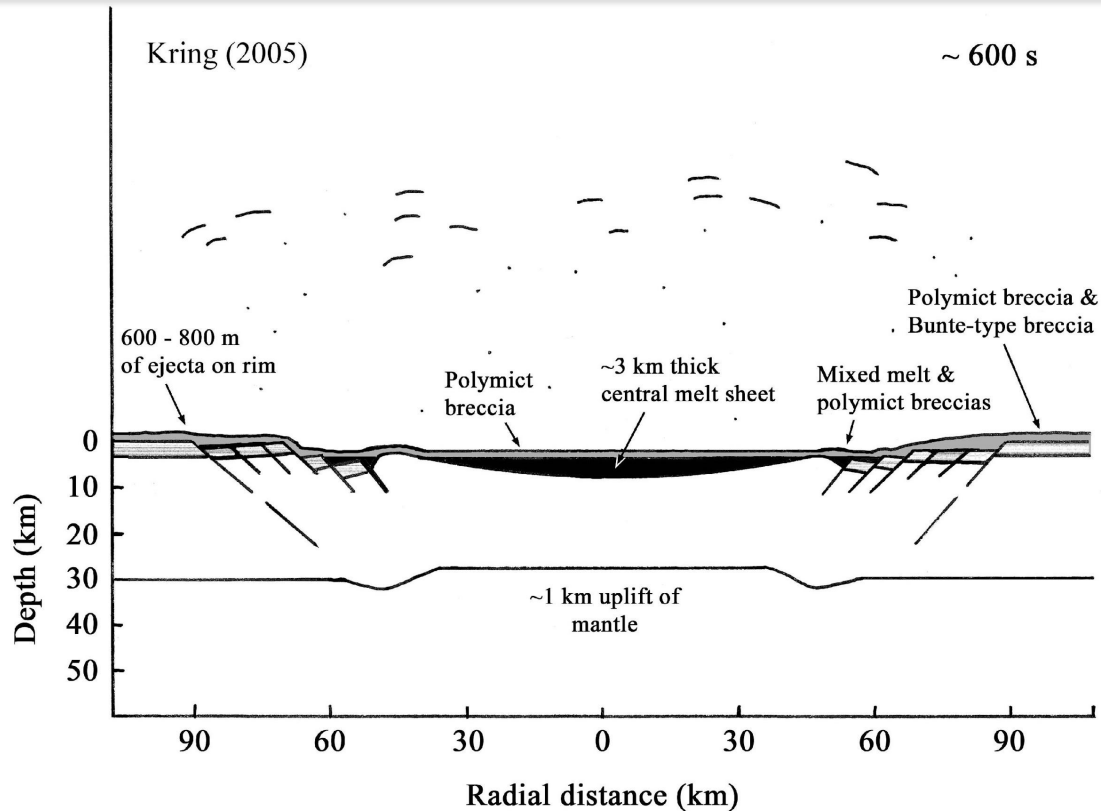


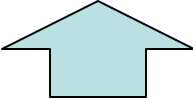
Fig. 6. Compilation of the pressure intervals over which certain shock effects are formed in quartz, olivine, graphite and calcite (modified after Stöfler and Langenhorst 1994 and Langenhorst and Deutsch 1998). The diagram is based on shock experiments with non-porous samples.

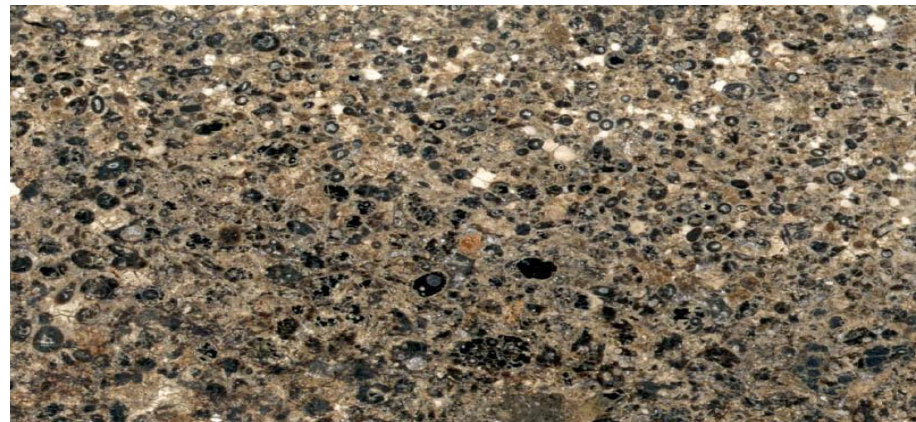
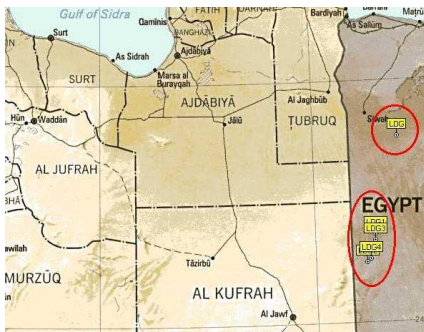
IMPACT MELT BODIES



Impact melting can profoundly affect crust evolution in the early stages of planet formation

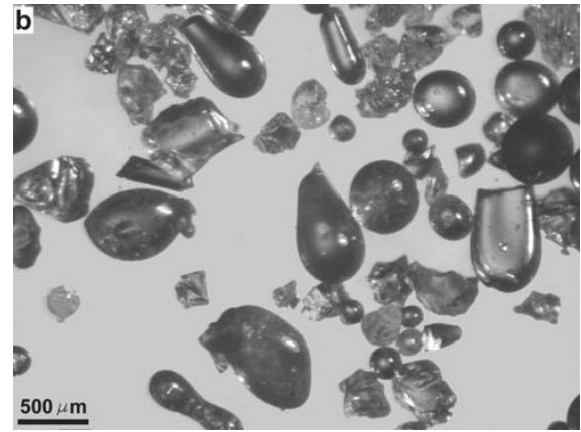
IMPACT MELT EVOLUTION?

- MELT PHYSICAL PROPERTIES evolution:
 - DENSITY
 - VISCOSITY
- 
- Melt composition, P, T
 - Redox conditions [$\text{Fe}^{3+}/(\text{Fe}^{2+} + \text{Fe}^{3+})$];
 - Water content.

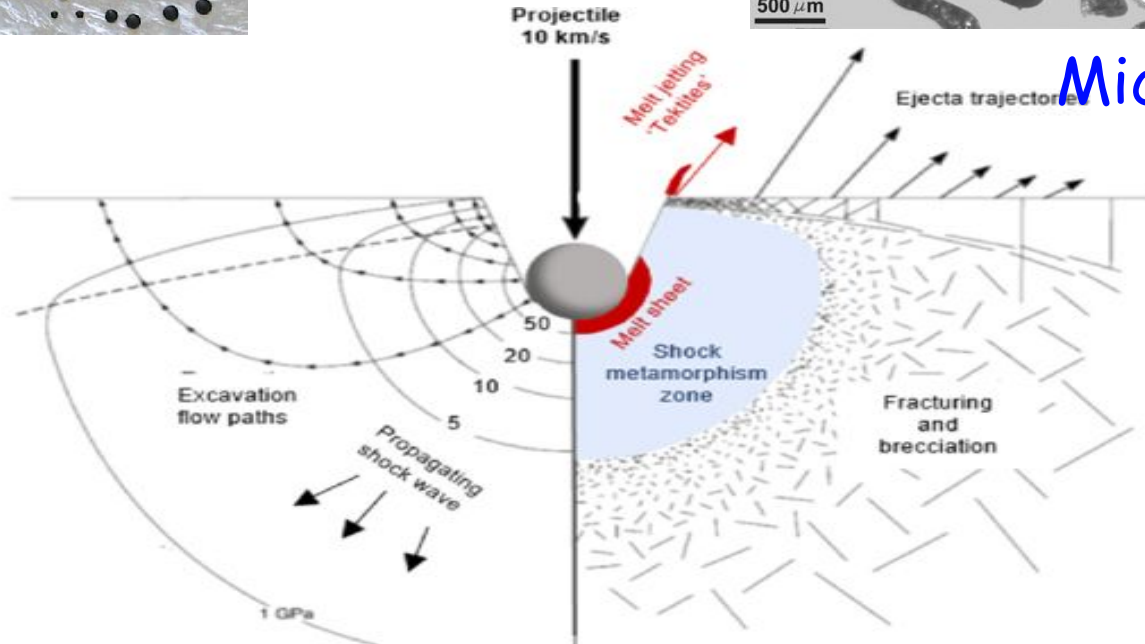




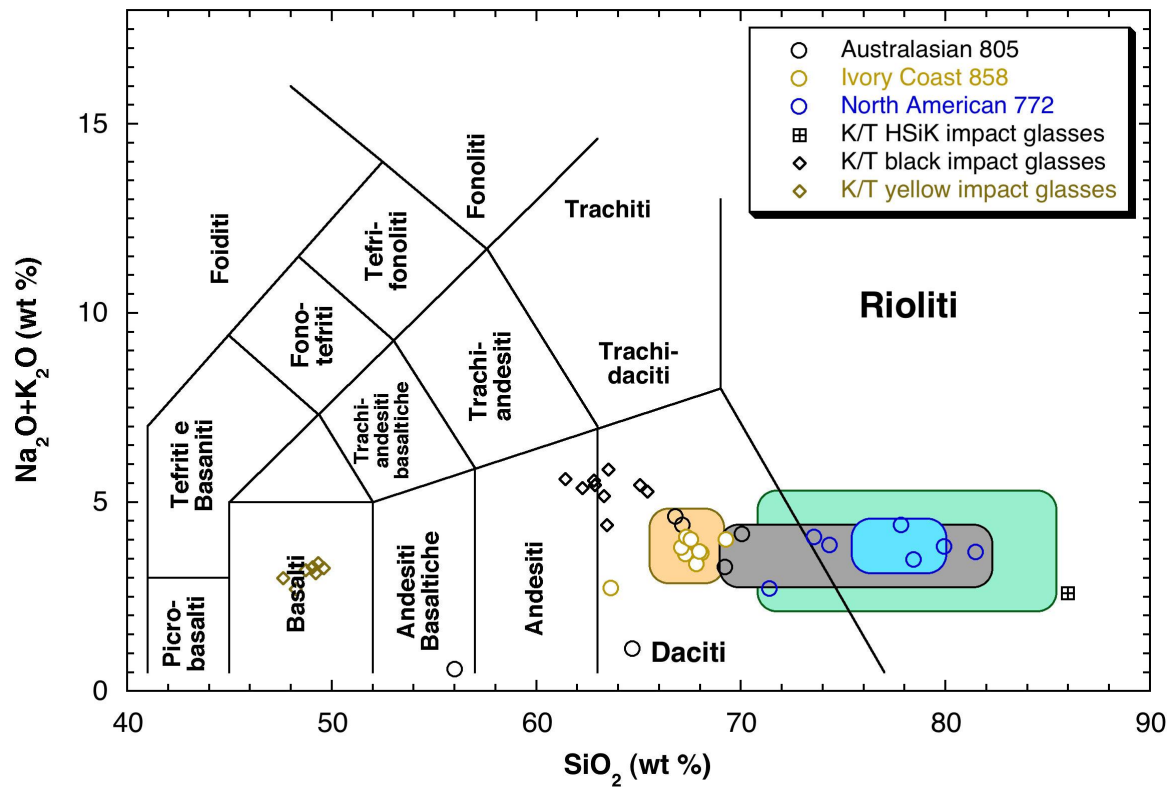
Tektites



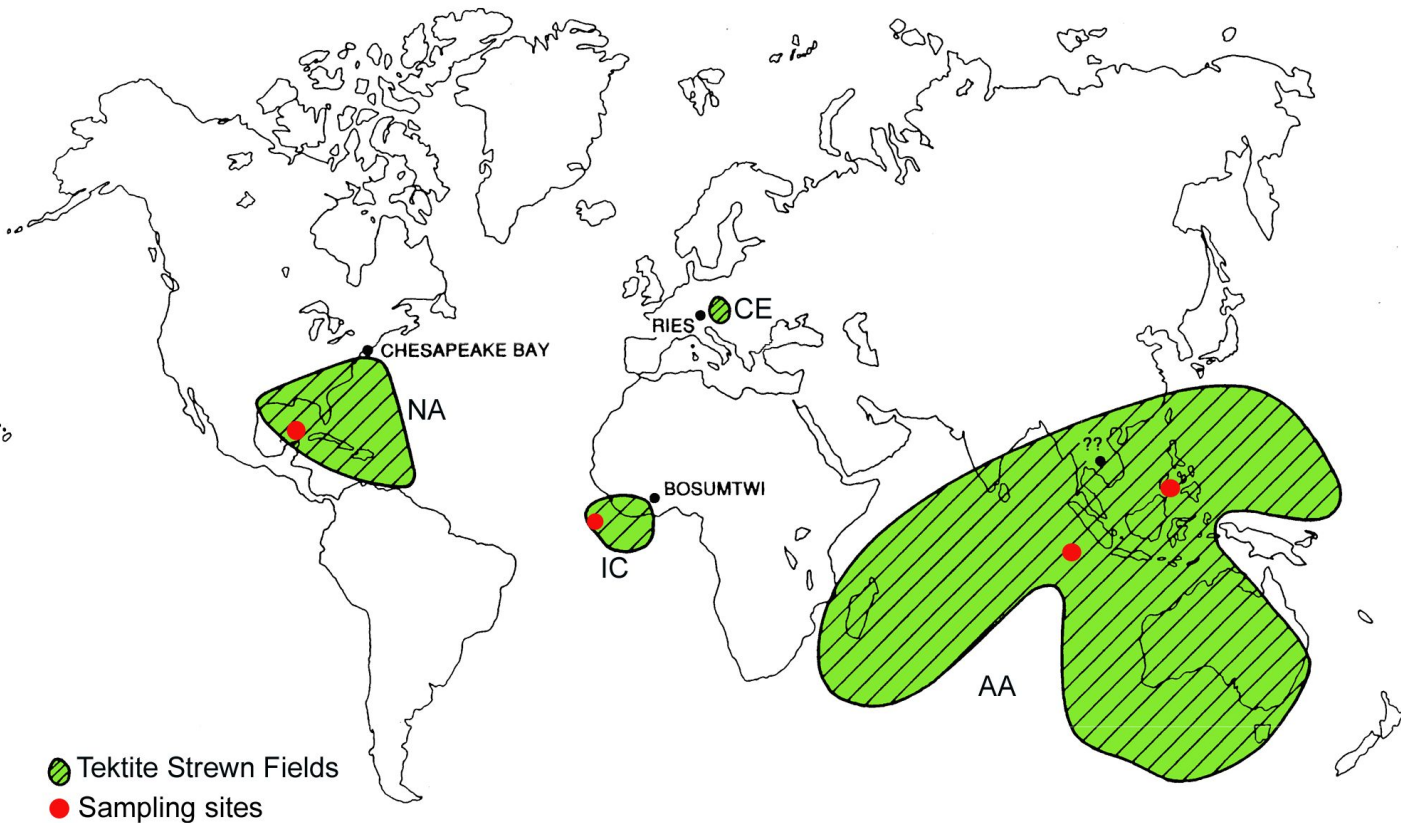
Microtektites



Microtektite and impact glass composition



Tektites

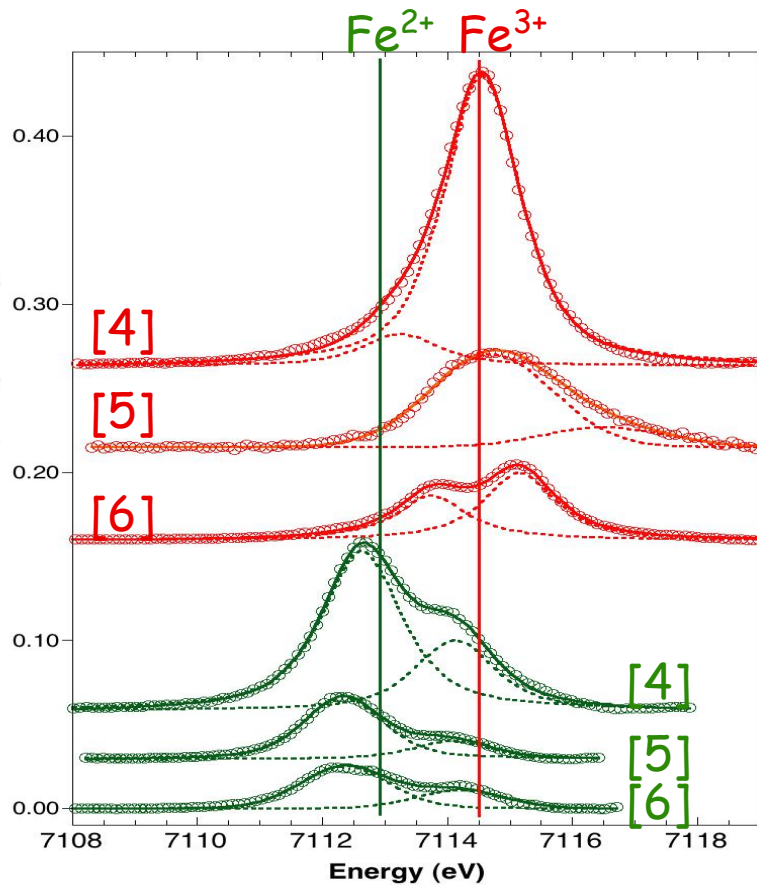
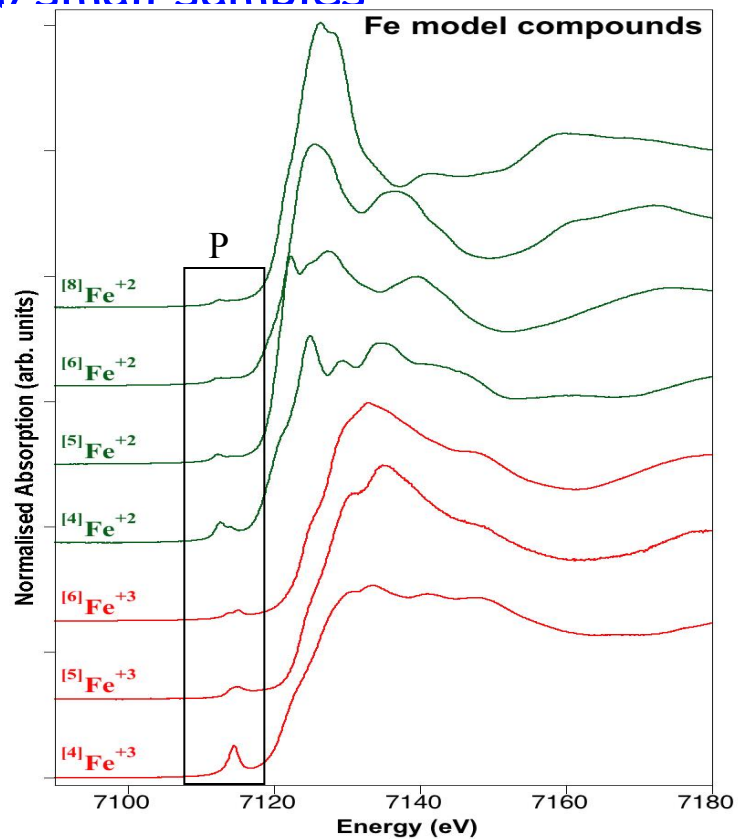


Effect of impact cratering on the planetary crust?

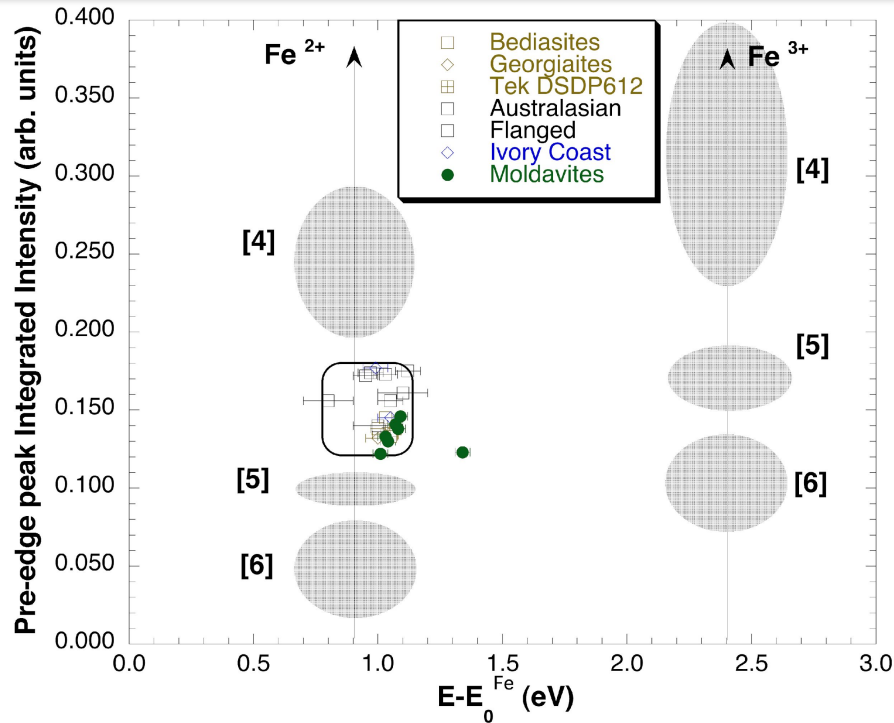
- Effect of impact cratering on melt:
 - melt redox: $Fe^{3+}/(Fe^{2+} + Fe^{3+})$;
 - melt water content.
- Transport processes of melt droplets

Analytical technique

Element selective technique able to analyse Fe redox also in diluted/small samples



Tektites



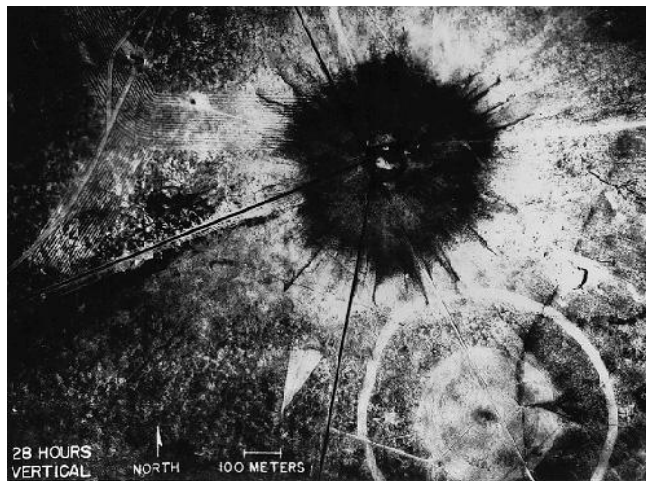
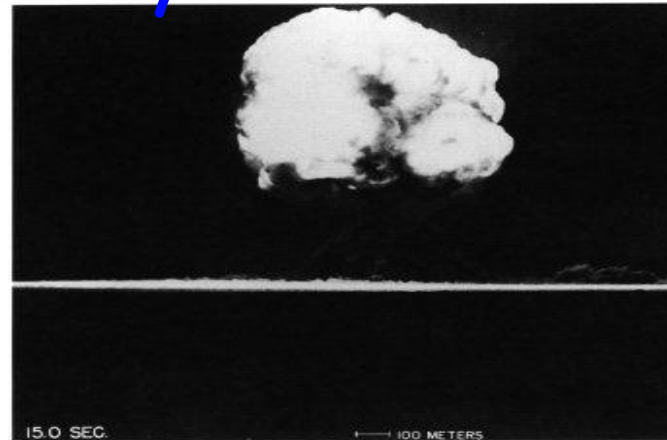
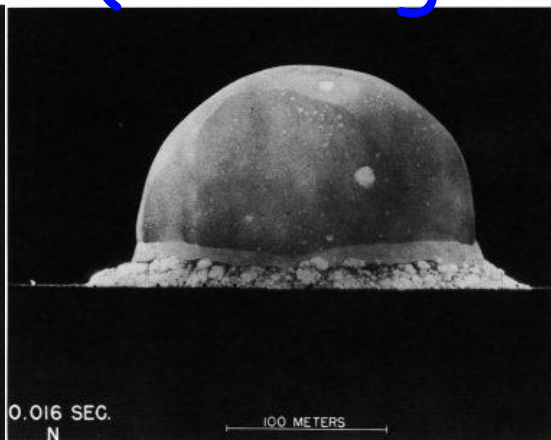
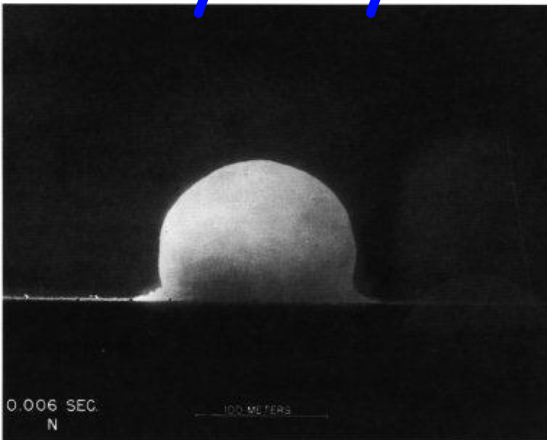
$Fe^{3+}/(Fe^{2+} + Fe^{3+}) \approx 0.05 \div 0.15 (\pm 0.03)$. Rarely up to 0.2

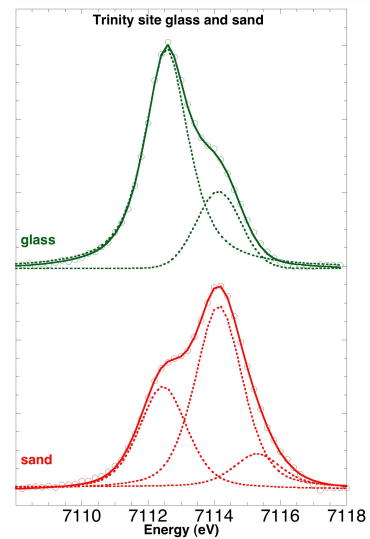
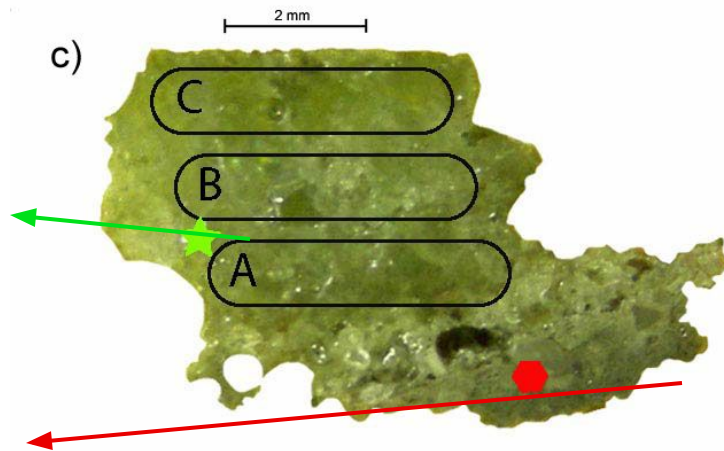
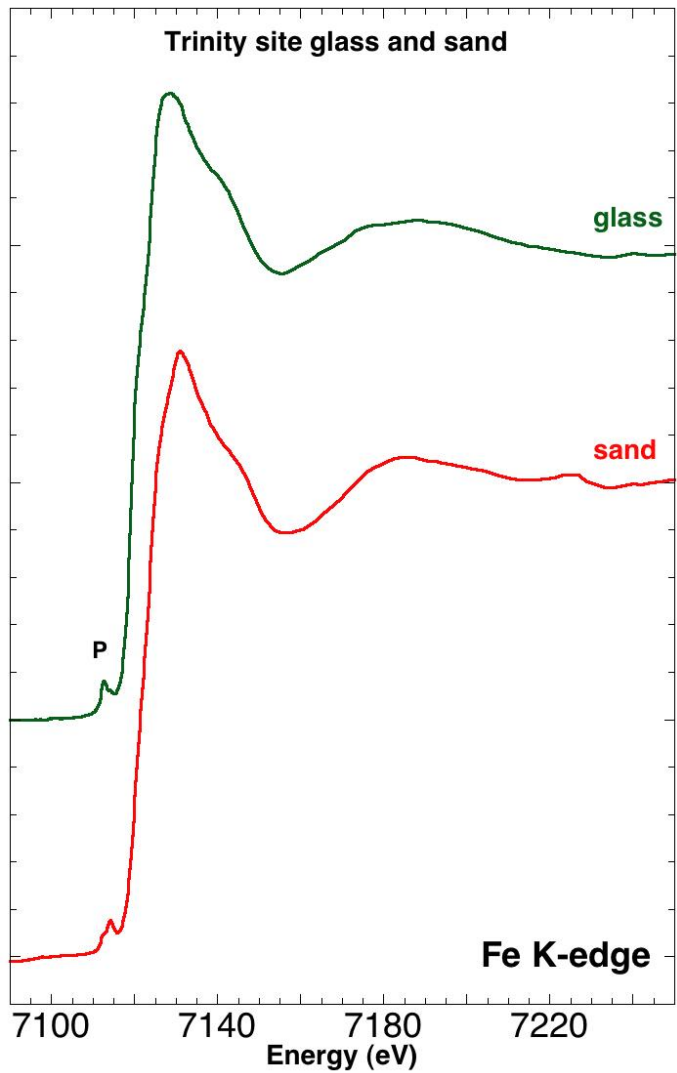
G. Giuli, G. Pratesi, E. Paris, C. Cipriani (2002) *Geochimica et Cosmochimica Acta*, **66**, 4347-4353.

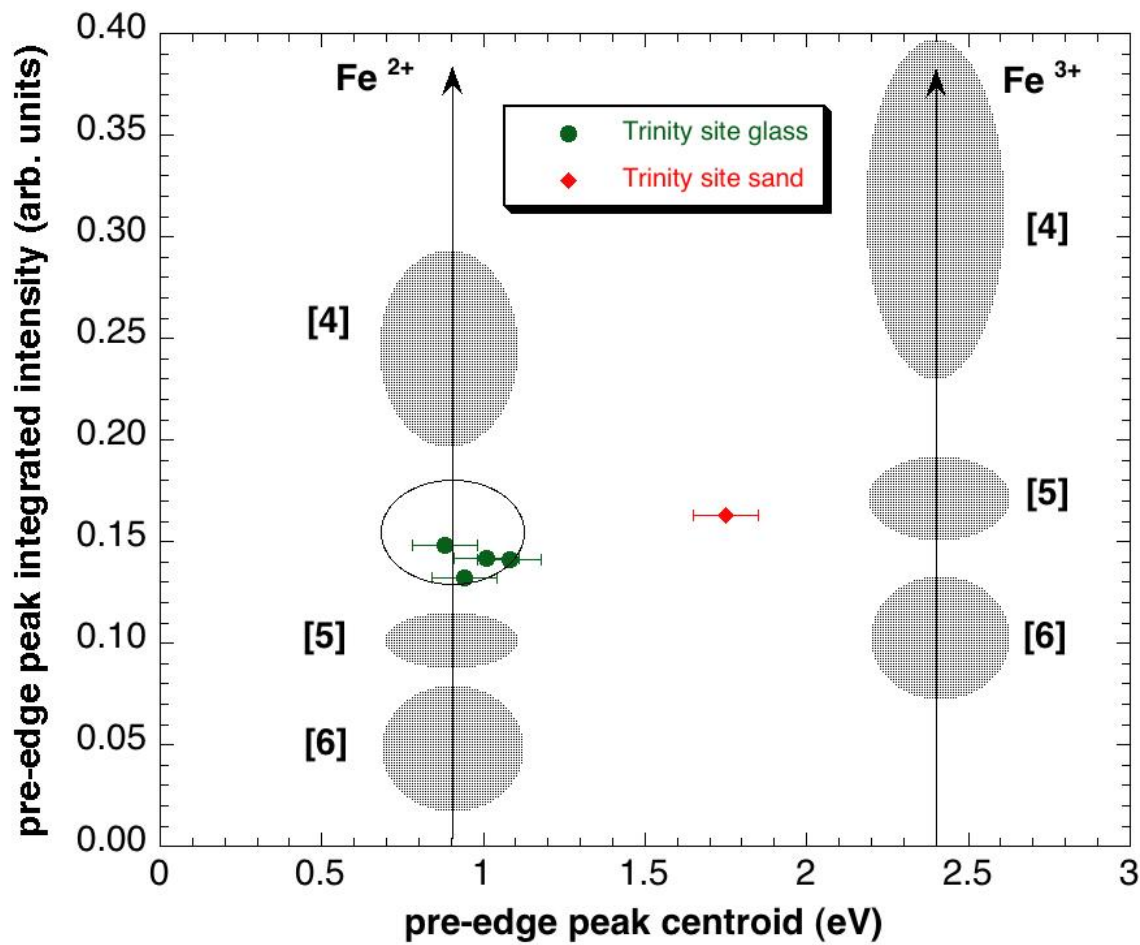
G. Giuli, G. Pratesi, S.G. Eeckhout, C. Koeberl, E. Paris, (2010a) *Geological Society of America Special Paper*, 465, 653-660.

G. Giuli, S.G. Eeckhout, M.R. Cicconi, C. Koeberl, G. Pratesi, E. Paris, (2010b) *Geological Society of America Special Paper*, 465, 645-651.

Trinity site (Alamogordo) July 16 1945

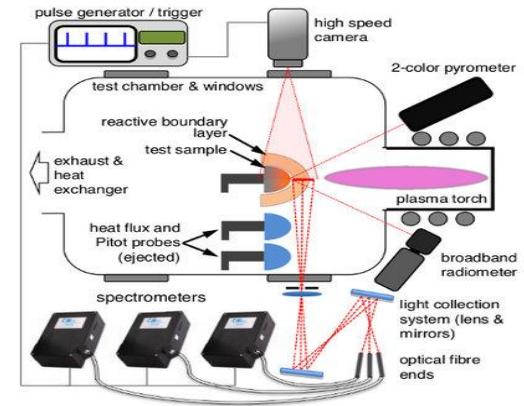
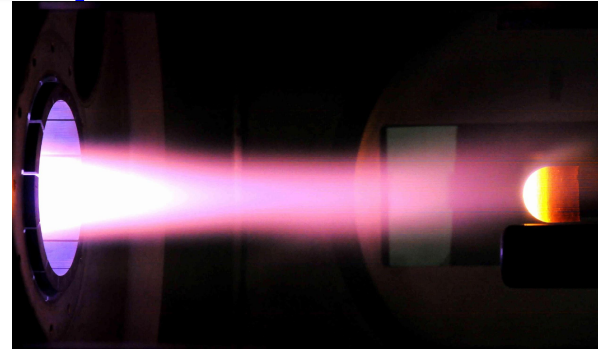
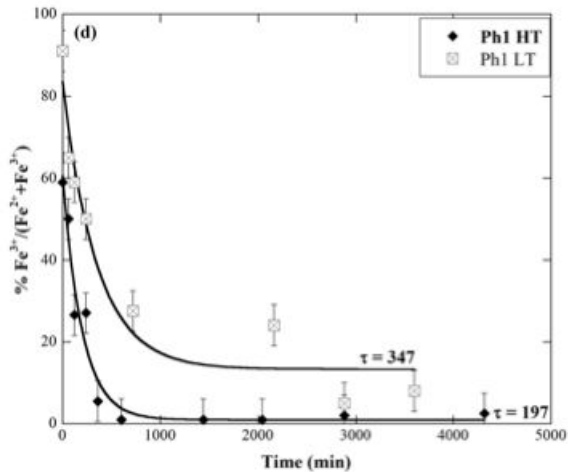






-Massive reduction in molten rock from the first atomic bomb test (Alamogordo, USA)

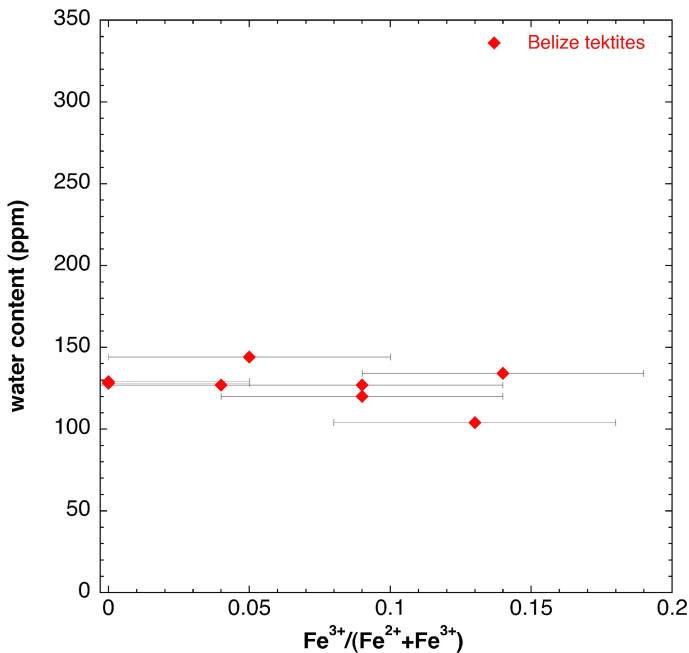
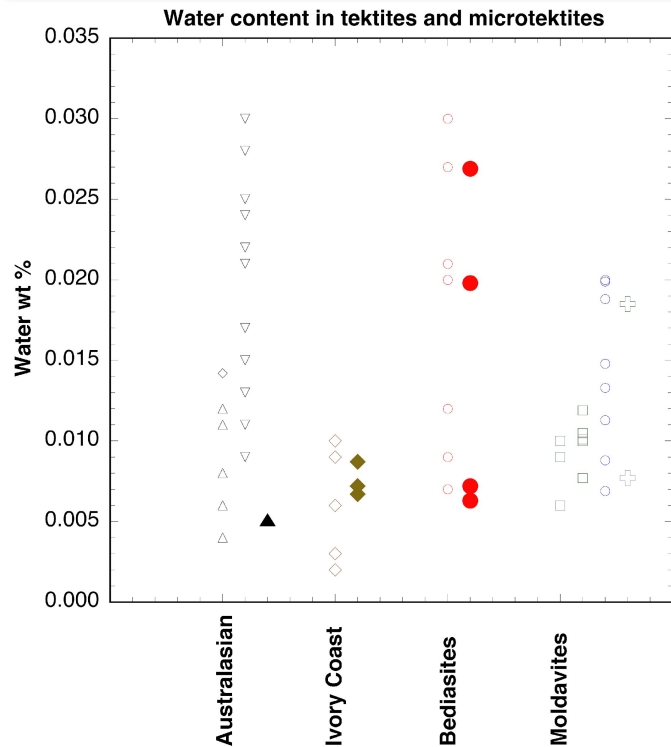
A matter of speed



Magnien, V., et al. (2006) *Journal of Nuclear Materials*, 352, 190-195.
Magnien, V., et al. (2008) *Geochimica et Cosmochimica Acta*, 72, 2157-2168.
Ciccioni et al. (2015) *Amer. Mineral*, 100, 1610-1619
Ciccioni et al. (2015) *Amer. Mineral*, 100, 1013-1016

$\text{Fe}^{3+}/(\text{Fe}^{2+}+\text{Fe}^{3+})$ 0.11-0.17 ± 0.05
Pittarello et al. (2019) *Icarus* 331, 170-178

Water content



Beran A., and Koeberl, C. (1997) *Meteoritics & Planetary Science*, 32, 211-216

Gilchrist J., Thorpe A.N., and Senftle F.E. (1969) *Journal of Geophysical Research*, 74, 1475-1483.

Giuli G., Cicconi M. R., Eeckhout S. G., Koeberl C., Glass B. P., Pratesi G., Cestelli-Guidi M., and Paris E. (2013). *Am. Mineral.* 98, 1930-1937

Giuli G., Cicconi M. R., Eeckhout S. G., Koeberl C., Glass B. P., Pratesi G., Cestelli-Guidi M., Carroll MR and Paris E. (2013). *Acta Minerl Petr.* 6, 777

Giuli, G., Cicconi, M.R., Stabile, P., Trapananti, A., Pratesi, G., Cestelli-Guidi, M., Koeberl, C., (2014). *LPI Contribution*, 1777, .2322.

Tektites:

-tektites have been homogeneously reduced from a presumably a wide range of Fe oxidation state in the target rock down to almost exclusively divalent.

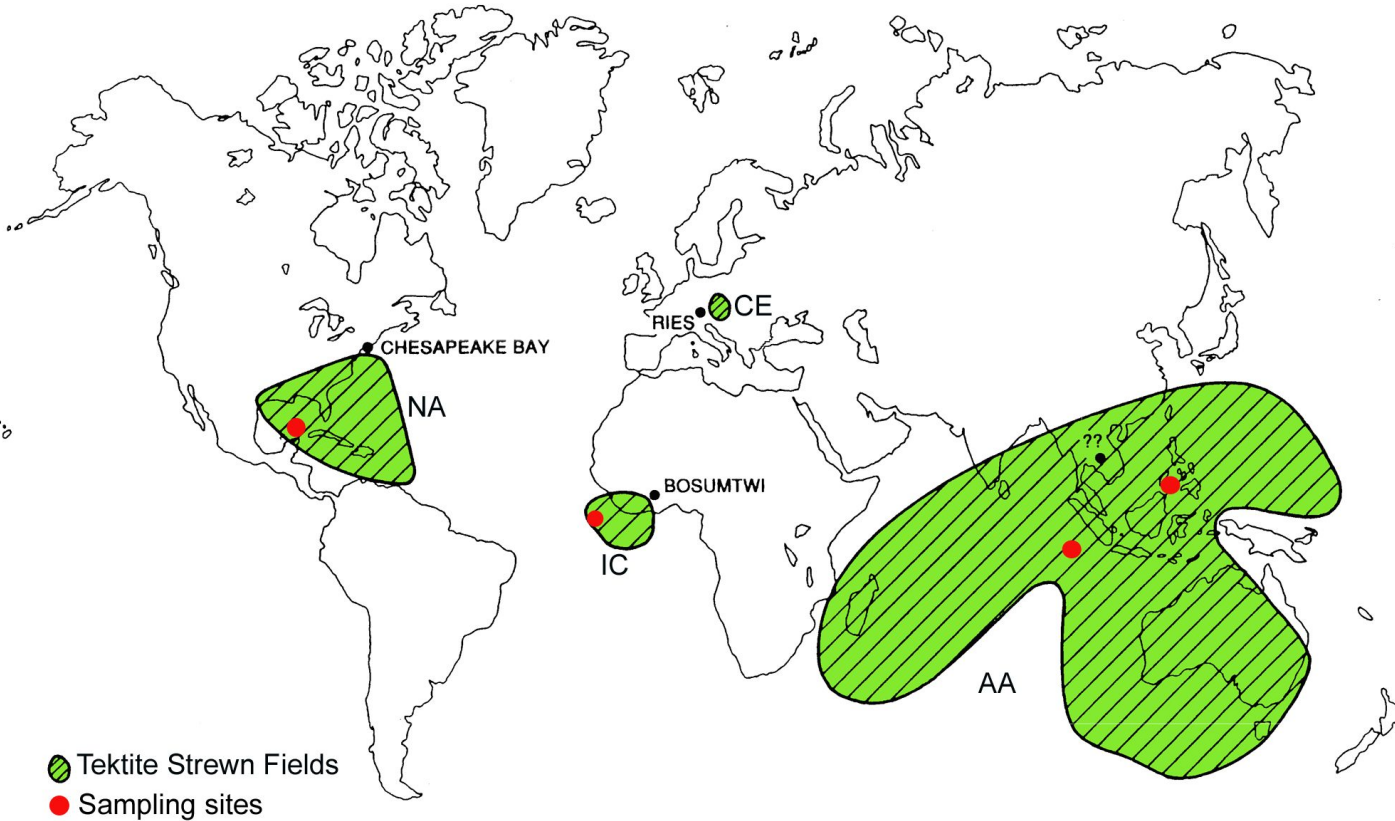
-No significant variations have been found according to impact age, target rock composition.

-G. Giuli, G. Pratesi, E. Paris, C. Cipriani (2002) *Geochimica et Cosmochimica Acta*, **66**, 4347-4353.

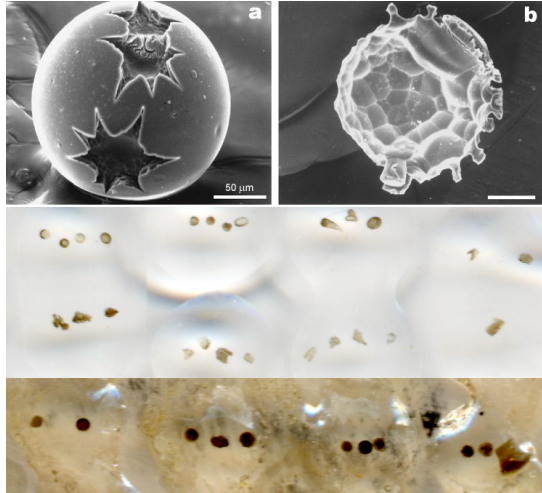
-G. Giuli, G. Pratesi, S.G. Eeckhout, C. Koeberl, E. Paris, (2010a) Large Meteorite Impacts and Planetary Evolution: *Geological Society of America Special Paper*, 465, 653-660.

-Giuli, G., Cicconi, M.R., Stabile, P., Trapananti, A., Pratesi, G., Cestelli-Guidi, M., Koeberl, C., (2014). LPI Contribution, **1777**, .2322.

Microtektites



Microtektites

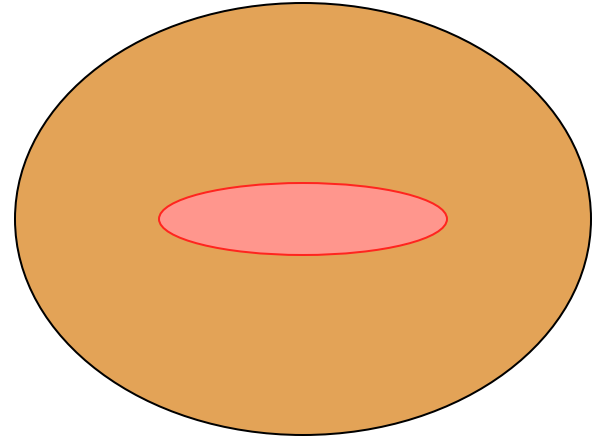


400-800 µm

Polished surfaces

Microprobe mount

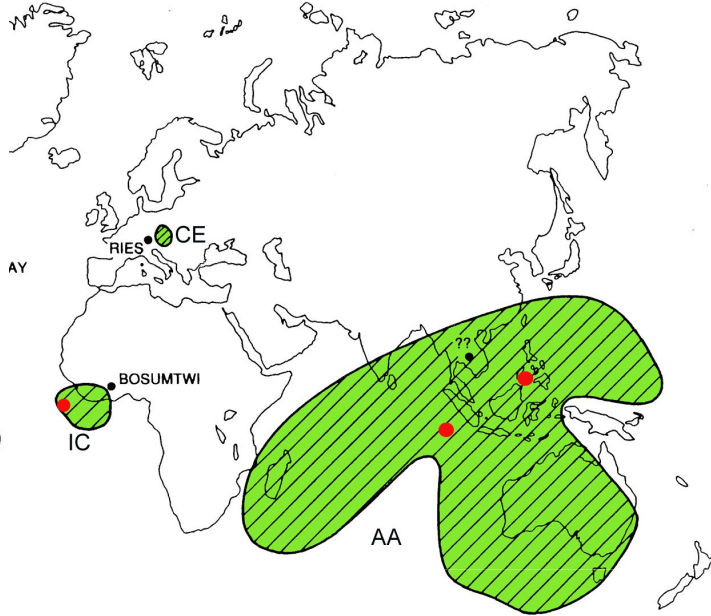
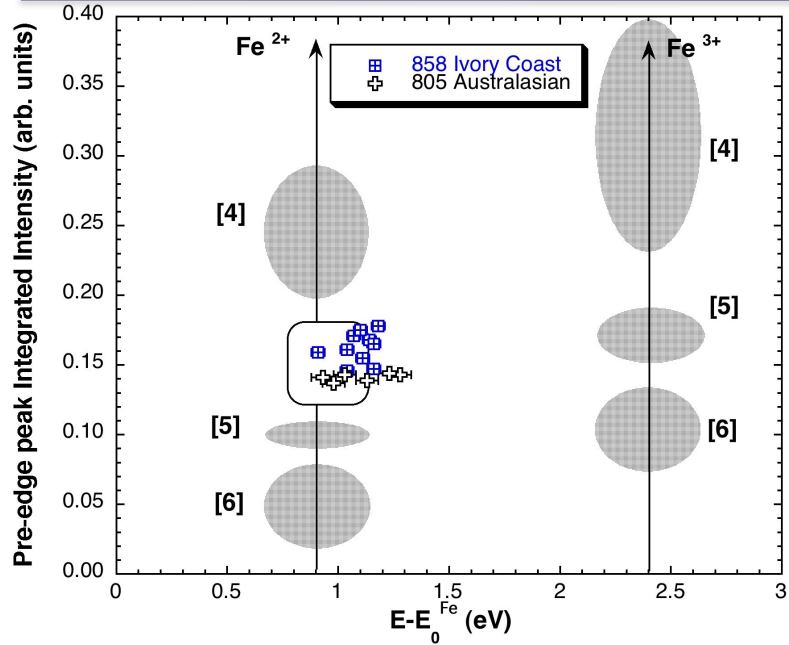
X-ray Beam



50 x 120 µm

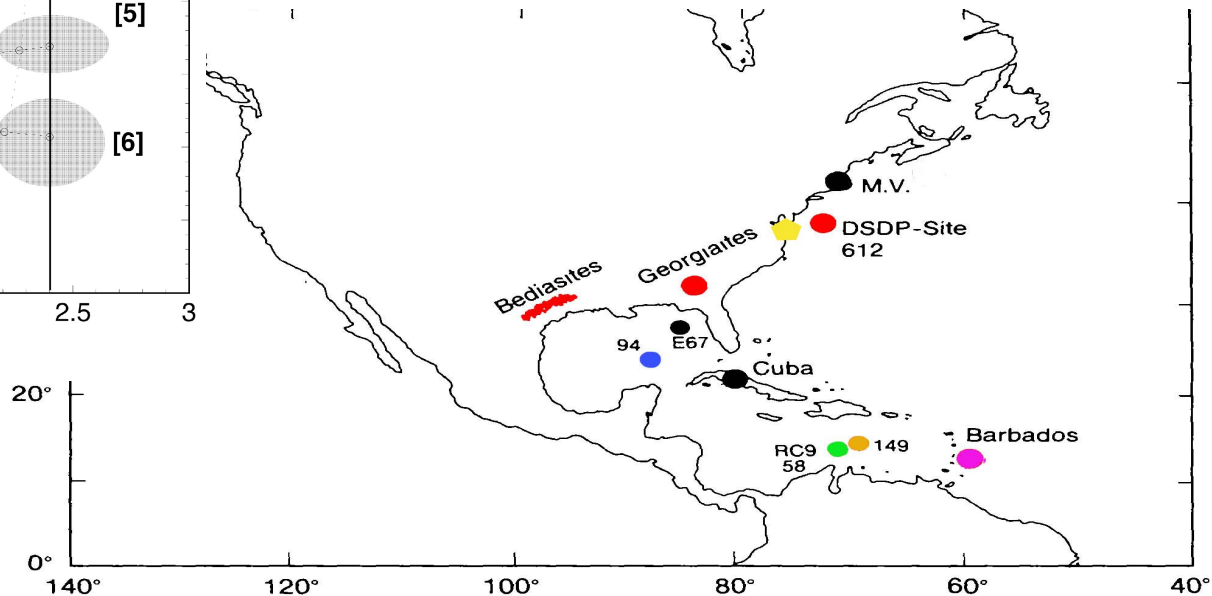
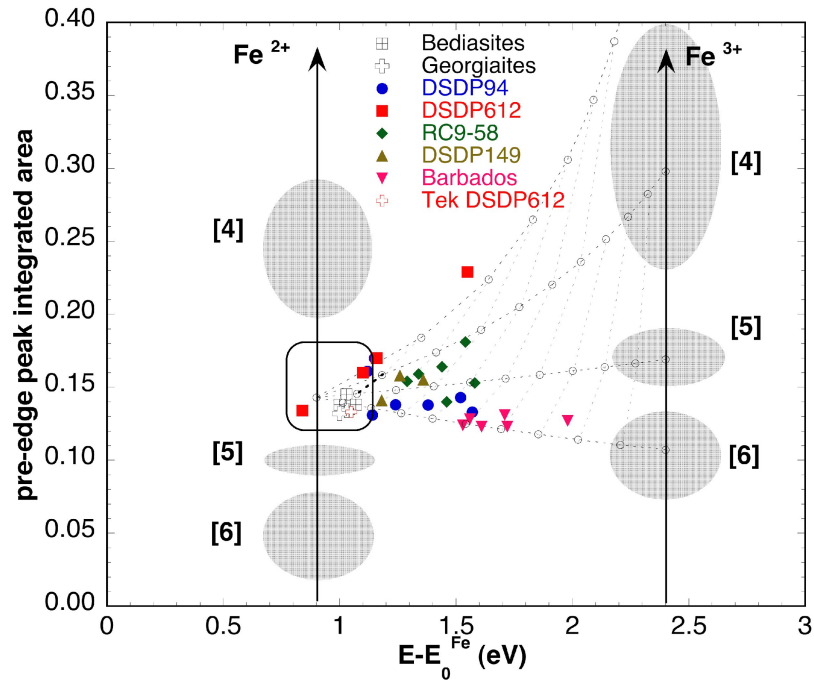
Sample at 45° from beam

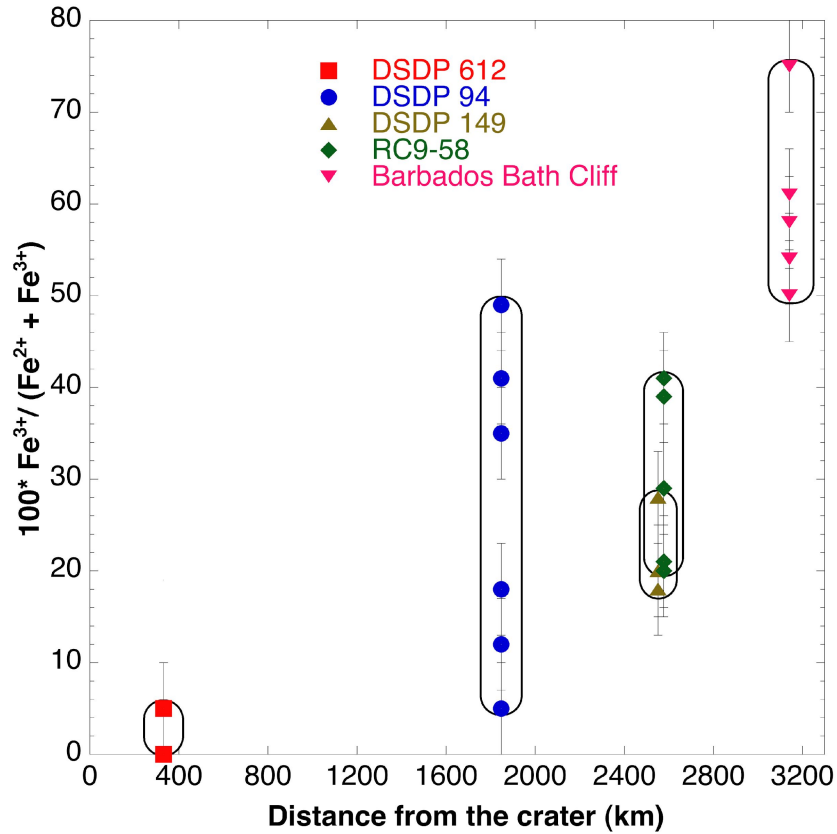
IC - AA Microtektites



- Tektite Strewn Fields
- Sampling sites

NA Microtektites





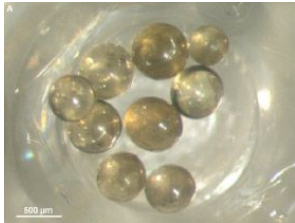
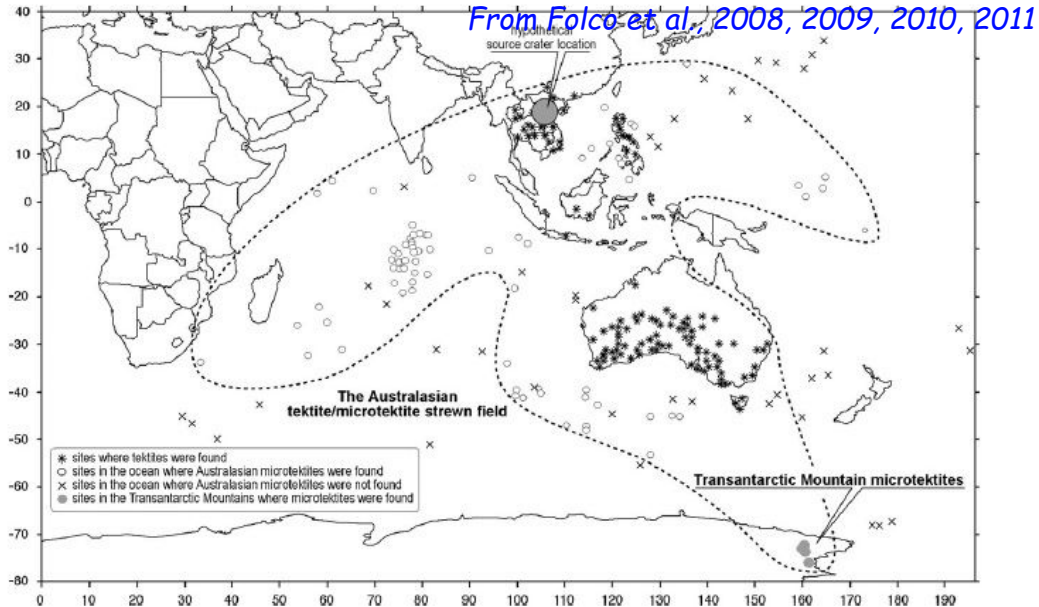
Distance from crater →

TAM Australasian Microtektites

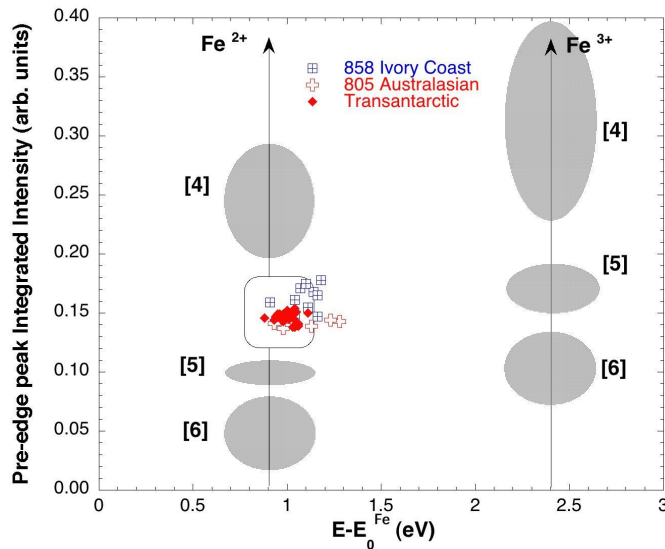
Comparable chemistry

Same age

Volatilisation trend vs. distance from crater



TAM Australasian Microtektites

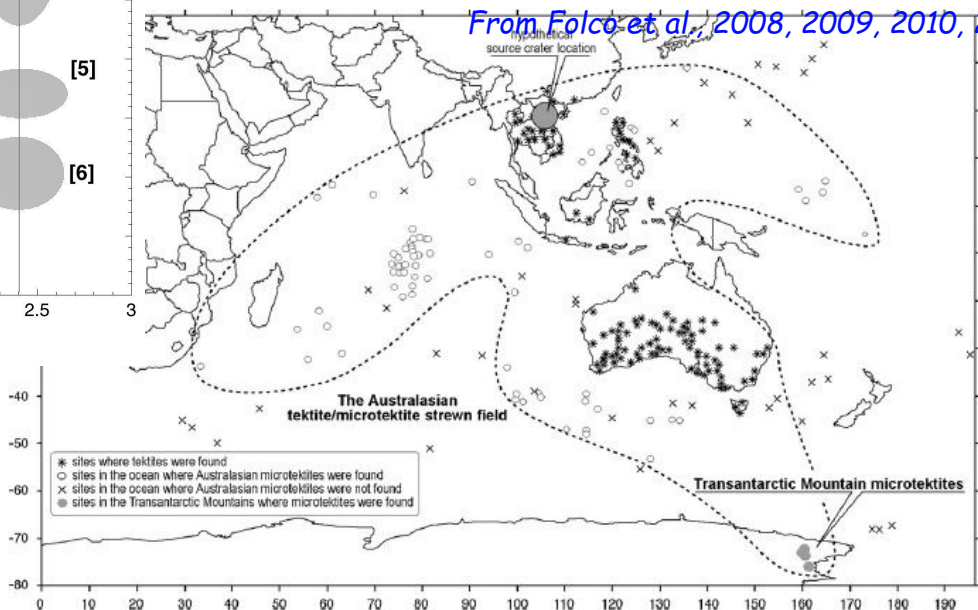


Comparable chemistry

Same age

Volatilisation trend vs. distance from crater

From Folco et al, 2008, 2009, 2010, 2011



Same Fe redox ratio as AU microtektites

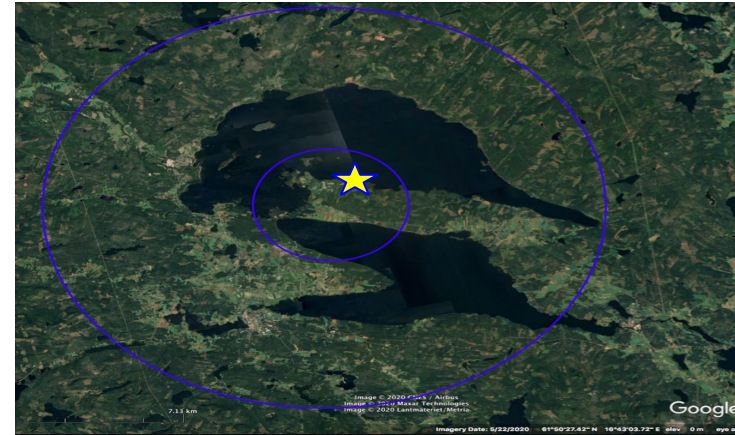
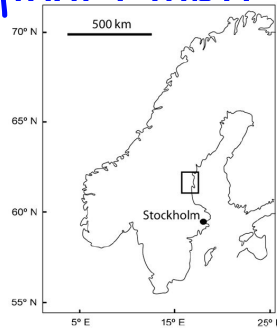
No oxidation during flight

Microtektites

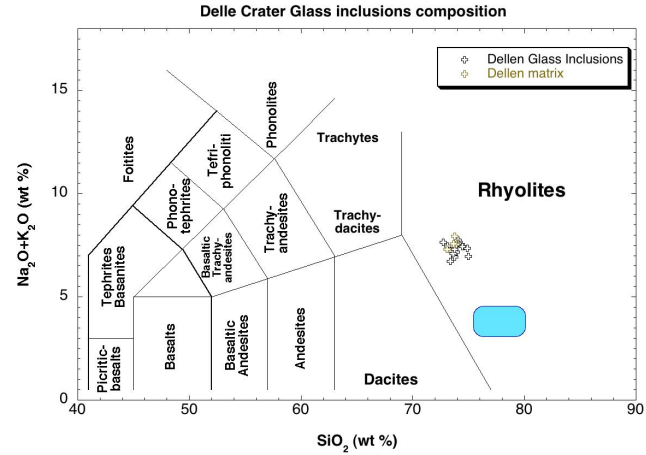
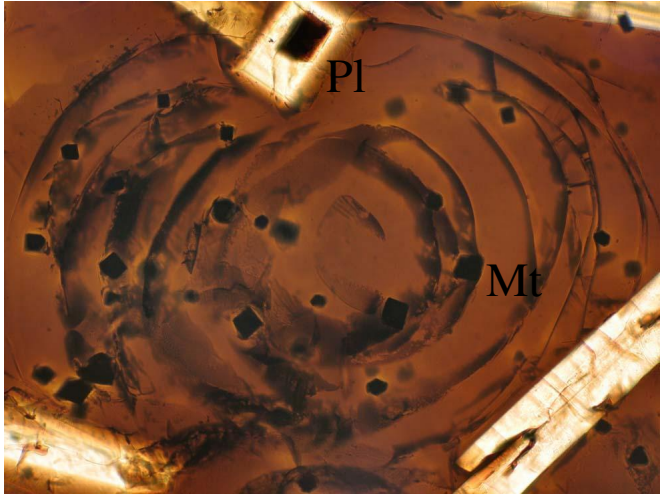
-
- In **Australasian** and **Ivory Coast** microtektites Fe is essentially divalent
- In **North-American** microtektites $\text{Fe}^{3+}/(\text{Fe}^{2+} + \text{Fe}^{3+})$ ratio ranges from 0 to 50% (correlation with distance)
- Possible differences in the formation/transport mechanisms with tektites?

Dellen (Sweden)

- rhyolite impact melt rocks (90-140 My old)
- \approx 200-500 m thickness of impact melt body
- glacially eroded
- melt evolution?



Dellen impact melt (Sweden)

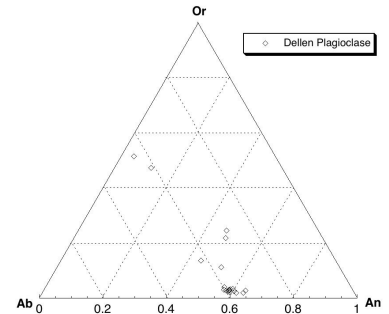
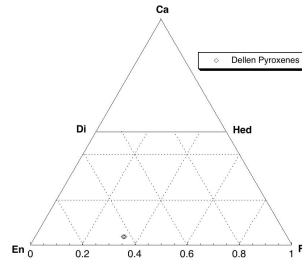
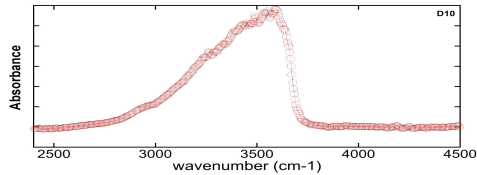


Px

- water = 1.47 ± 0.3 wt%

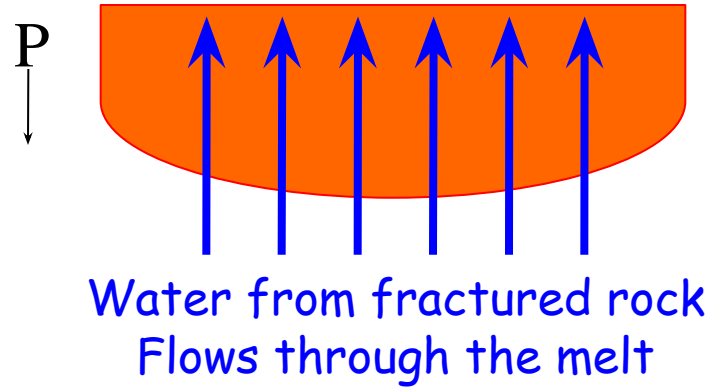
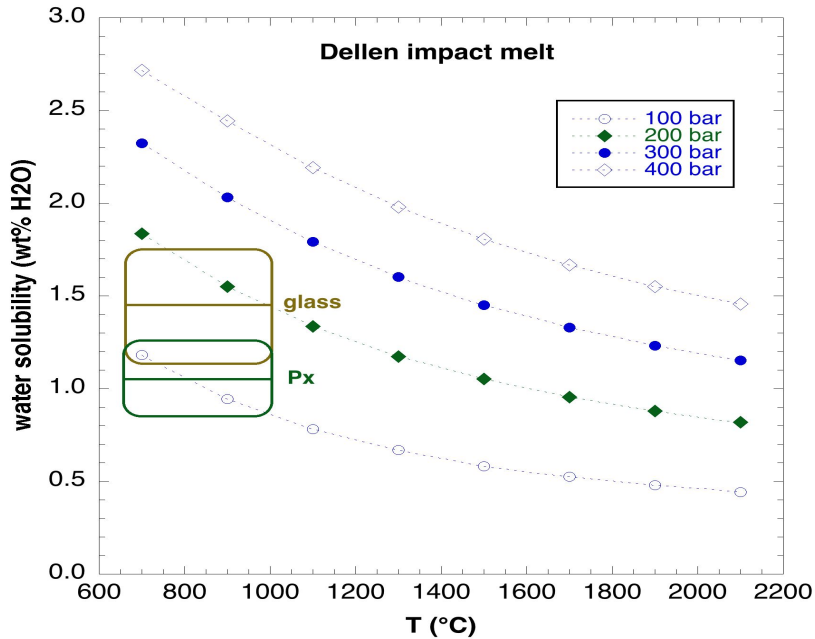
- P, T?

- origin?



Magma water content from pyroxene/glass

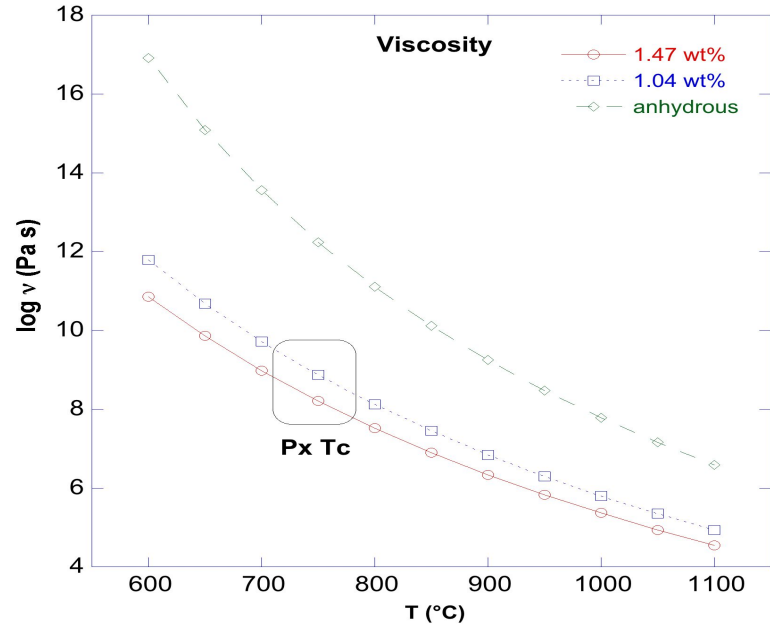
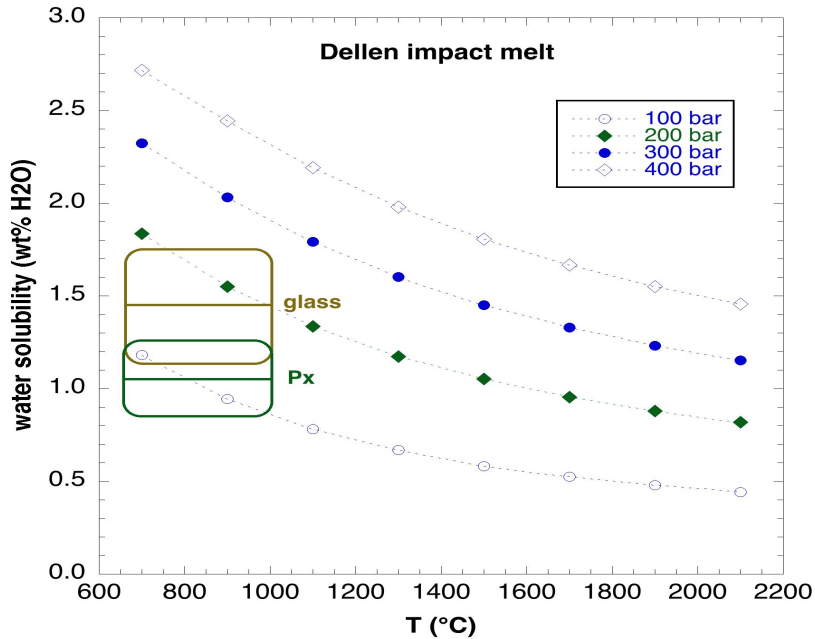
Using the partition coefficient of Tenner et al (2009) we calculated a value of 1.04 ± 0.21 wt% of H_2O^T for the melt in equilibrium with pyroxene.



H₂O solubility calculated according to Papale (1999) model; viscosity according to GRD (2006) model

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Dellen impact melt (Sweden)

- Dellen rhyolitic glass contains 1.47 ± 0.3 wt% water
- OH content of pyroxenes = 98 ppm OH

Resulting in melt water content of $1.04 \text{ wt\%} \pm 0.21$

- we interpret this water to originate from volatiles released by the shocked basement, diffusing through the impact melt
- water content of impact melt is essential in order to correctly model melt viscosity, density, and evolution

Conclusions

- impact cratering is a geologic process which has been very intense in the first billion y
- affect planetary surfaces both morphologically and chemically
- Impact melt formation can potentially affect crustal evolution
- water content of impact melt is essential in order to correctly model melt viscosity, density, and evolution
- Important implications to model the role of impact melts on water mobilization in planetary crusts

Thank you for your attention!



Carancas, Perú, 2007

Thanks to:

G. Pratesi, M. Corazza, C. Cipriani, E. Paris, C. Koeberl,
S. G. Eeckhout, M.R. Cicconi, M.R. Carroll, B. Glass, P. Stabile,
M. Cestelli Guidi, L. Folco, R. Skala, L. Dzikova, M. Trnka,
S. Caporali, F. Arzilli, W. Marinucci, K-U Hess, D.B. Dingwell,
S. Nazzareni

B. Glass

Uwe Reimold

Philippe Claeys

discussions

Museo Storia Naturale Firenze

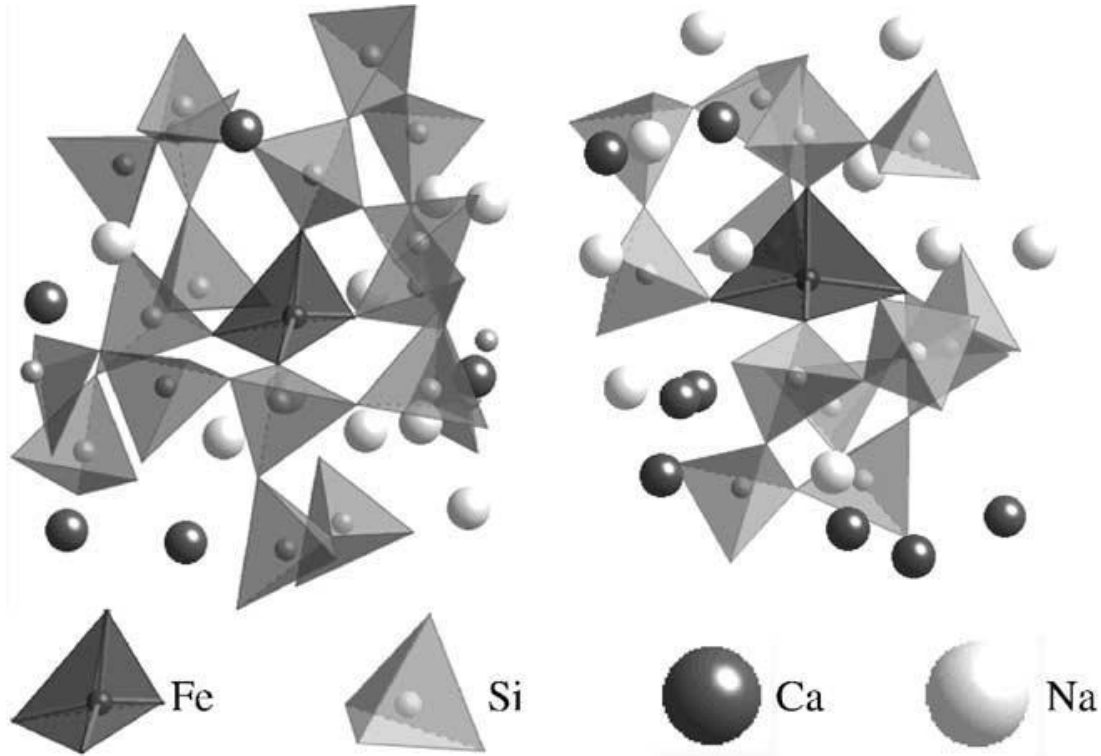
NHM Vienna

Museo Nazionale dell'Antartide (Siena)

DSDP

samples

Structure of silicate glasses



Structure of silicate glasses

Fe^{2+}

[4]

$\langle \text{Fe-O} \rangle \sim 2.00 \text{ \AA}$

[5]

$\langle \text{Fe-O} \rangle \sim 2.07 \text{ \AA}$

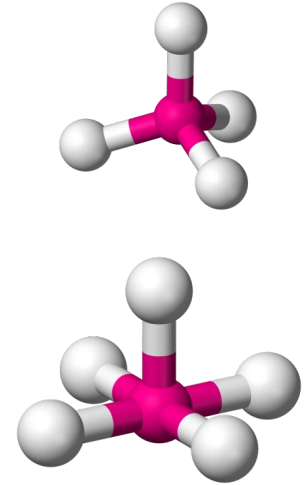
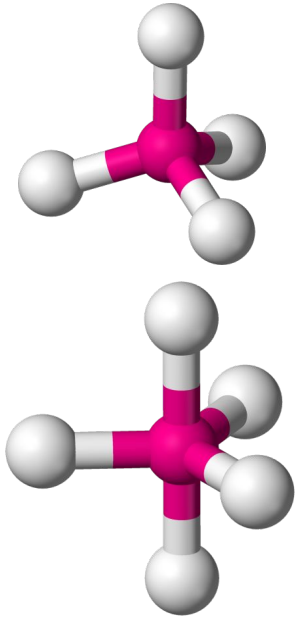
Fe^{3+}

[4]

$\langle \text{Fe-O} \rangle \sim 1.85 \text{ \AA}$

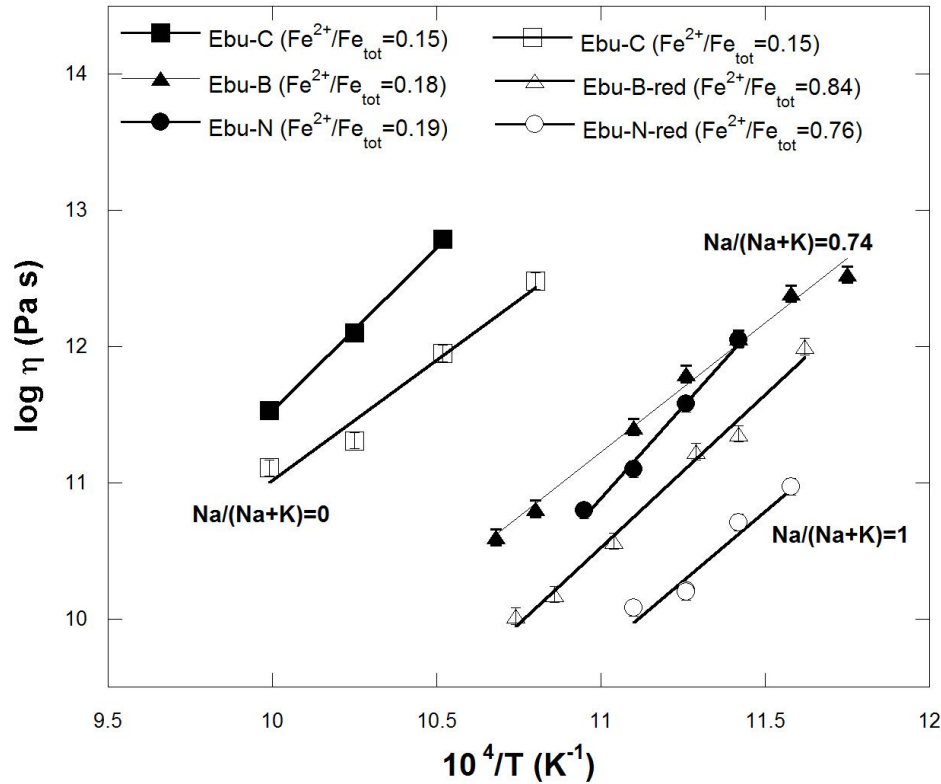
[5]

$\langle \text{Fe-O} \rangle \sim 1.94 \text{ \AA}$

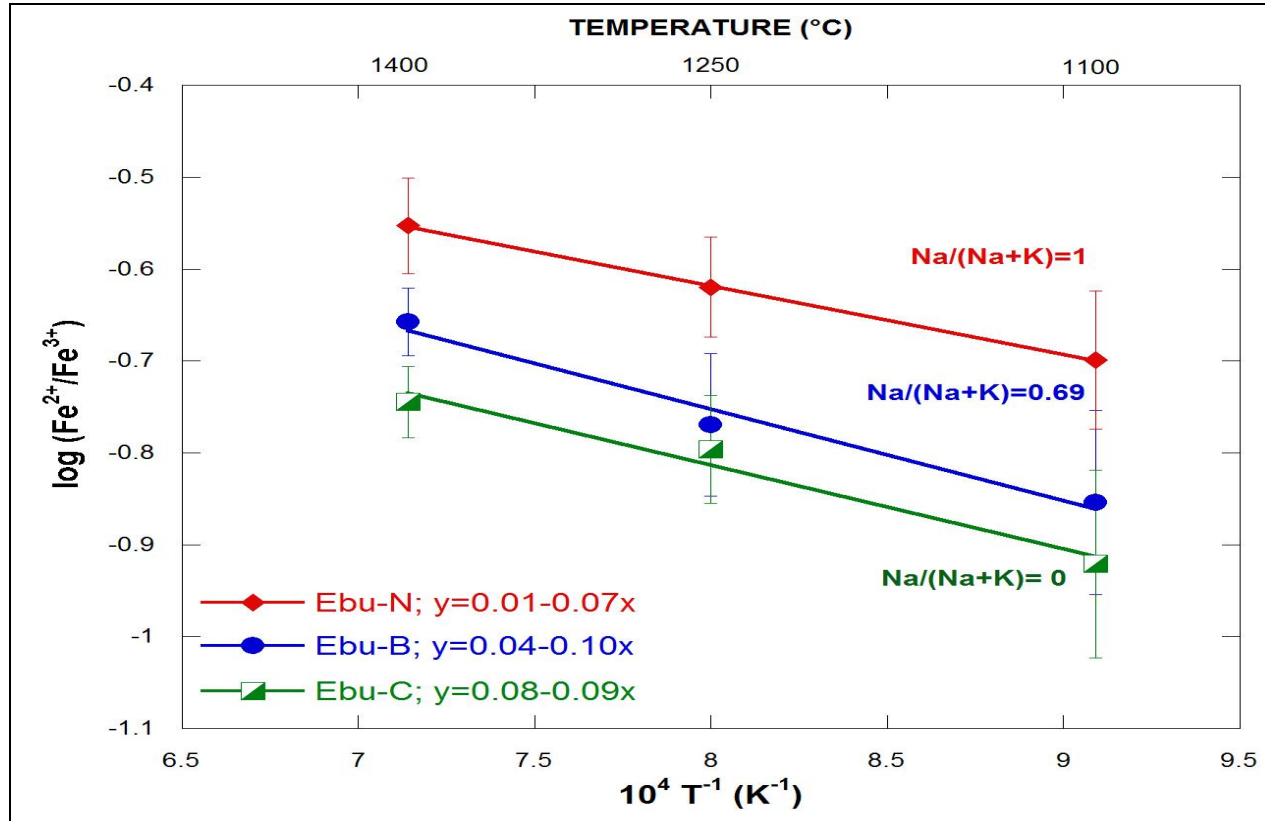


Density, compressibility, bond strength, polymerization

Alkalis and iron redox influence on viscosity

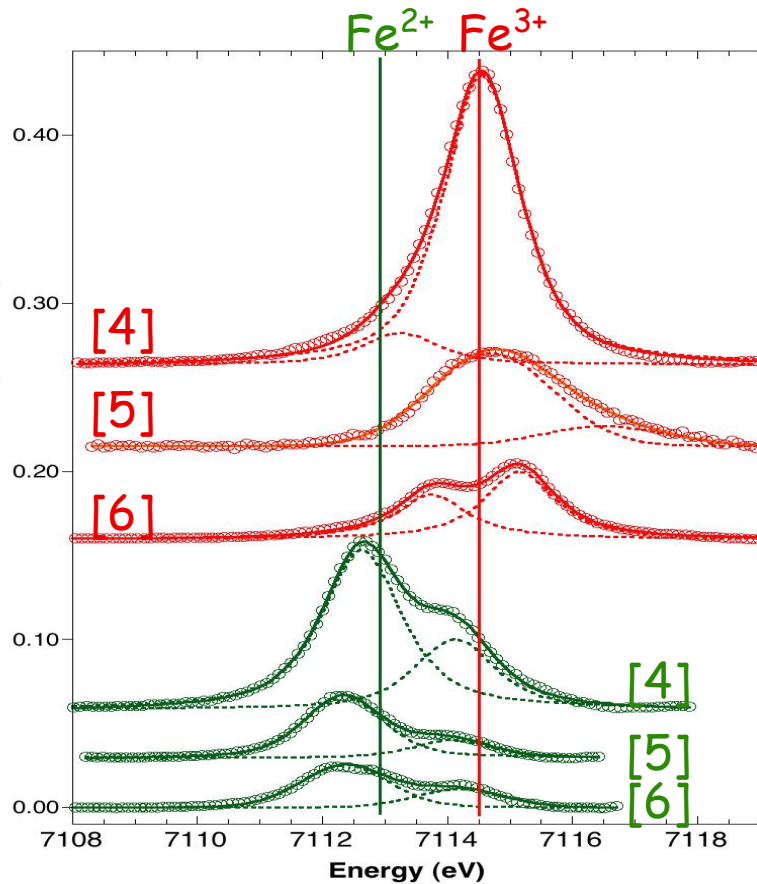
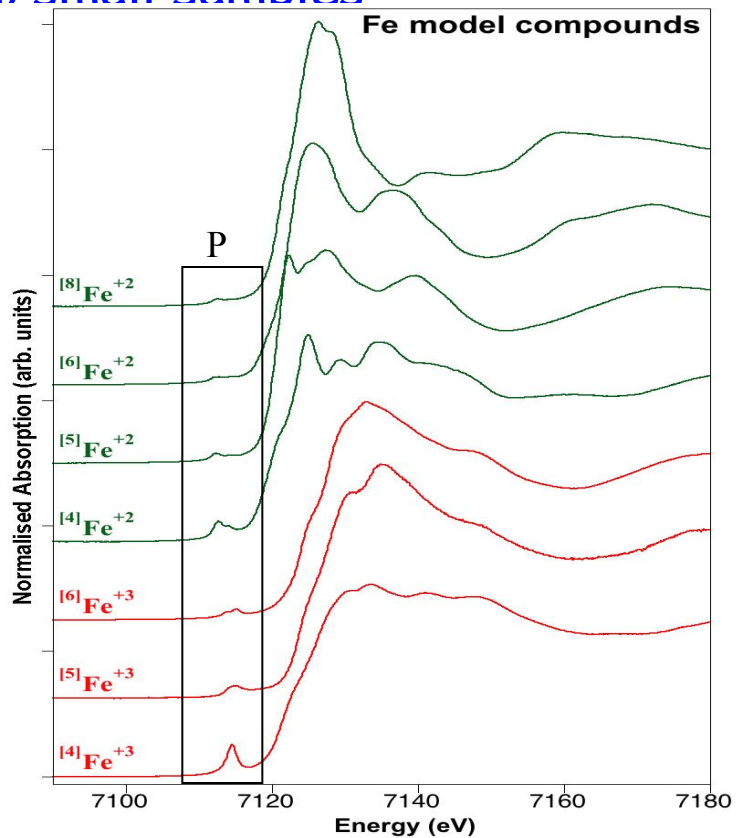


T and alkali influence on Fe redox



Analytical technique

Element selective technique able to analyse Fe redox also in diluted/small samples



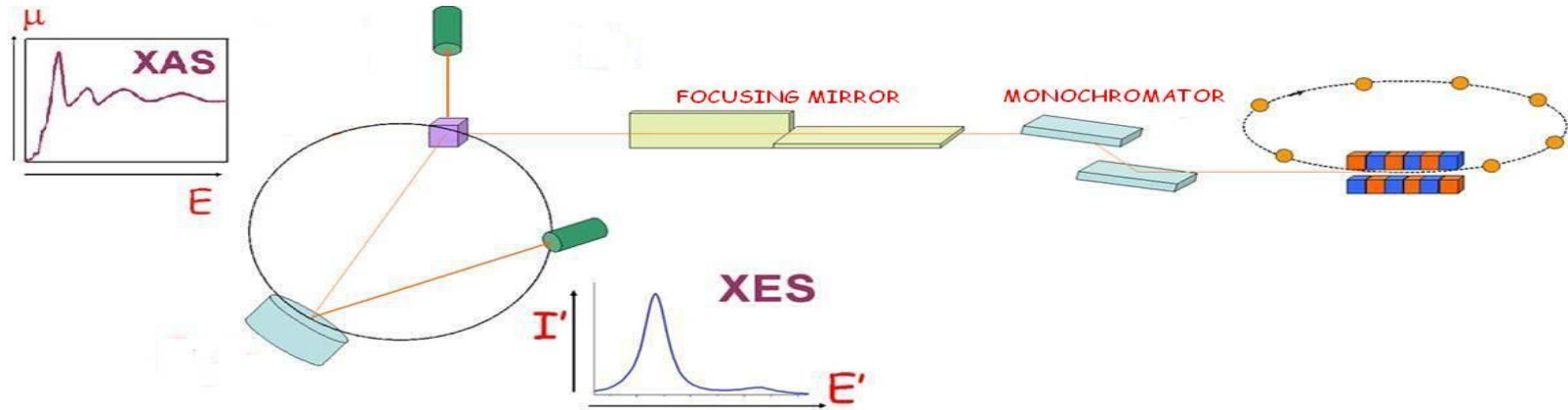
X-ray Absorption Spectroscopy

Element selective

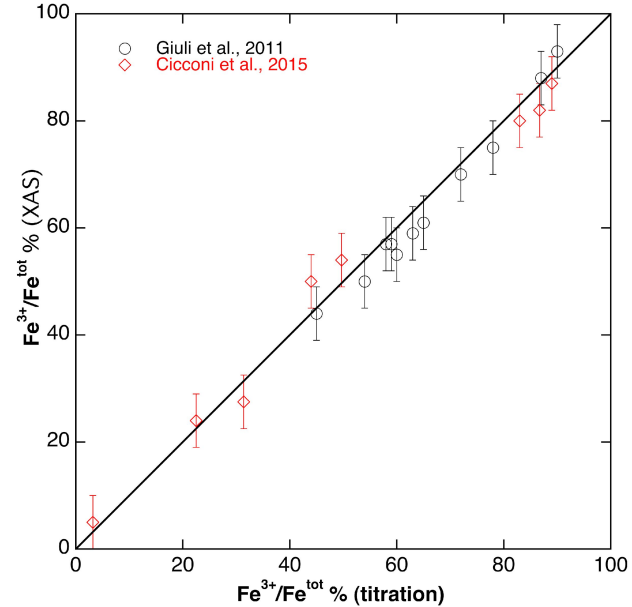
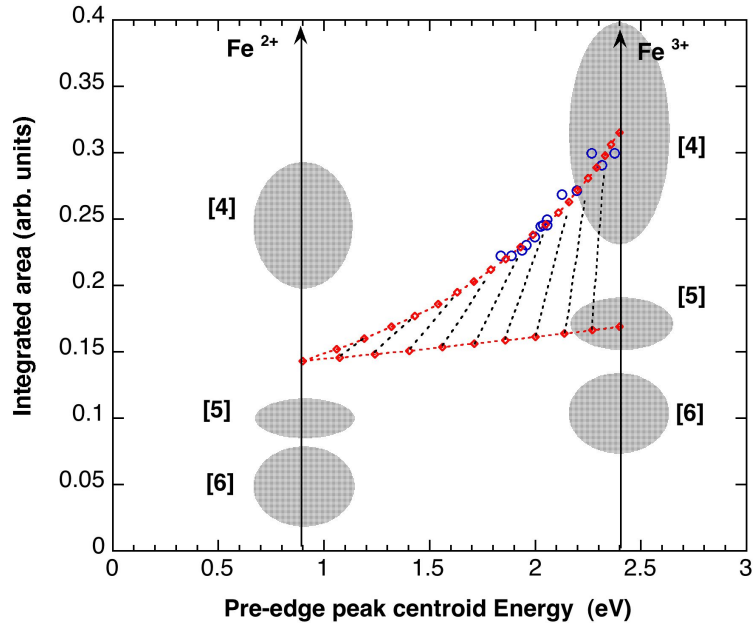
Provide oxidation state and local structure

Down to very low absorber concentration

Micro-probe



Calibration of the $\text{Fe}^{3+}/(\text{Fe}^{2+} + \text{Fe}^{3+})$ ratio

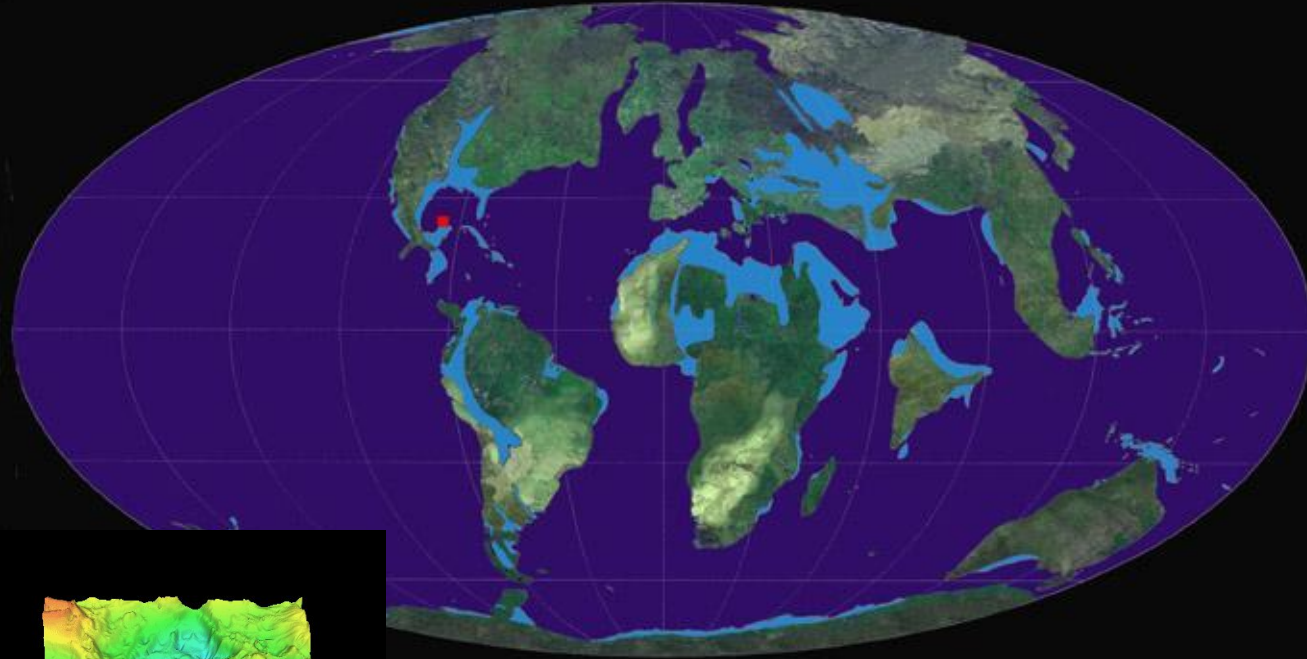


$\text{Fe}^{3+}/(\text{Fe}^{2+} + \text{Fe}^{3+})$ ratio ± 0.05

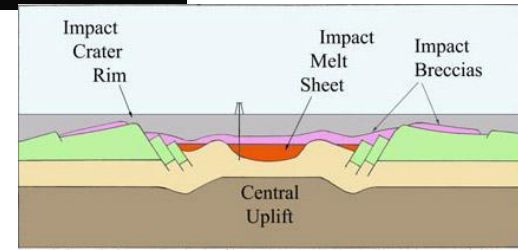
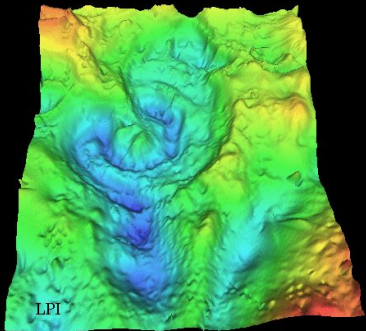
Giuli et al. (2011) *Amer. Mineral*, 96, 631-636

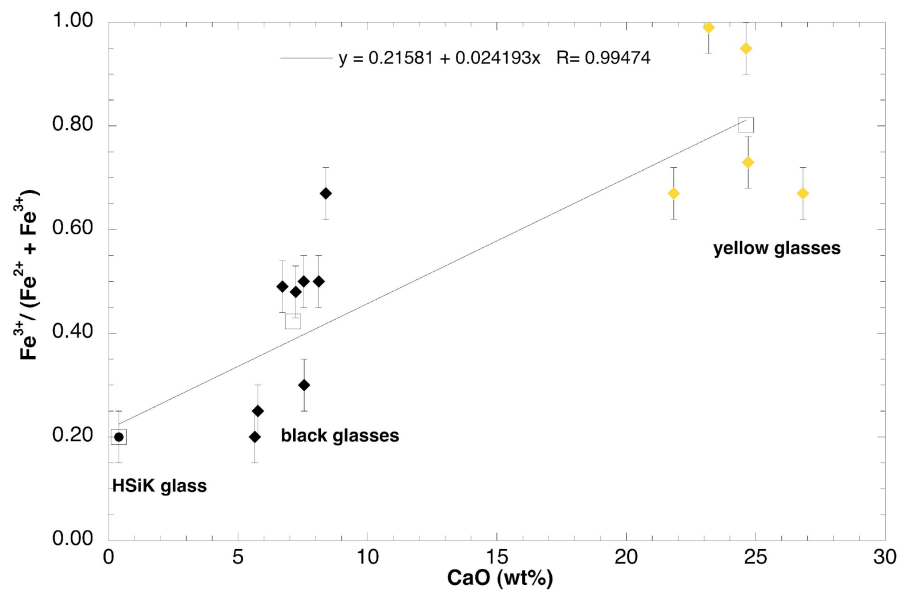
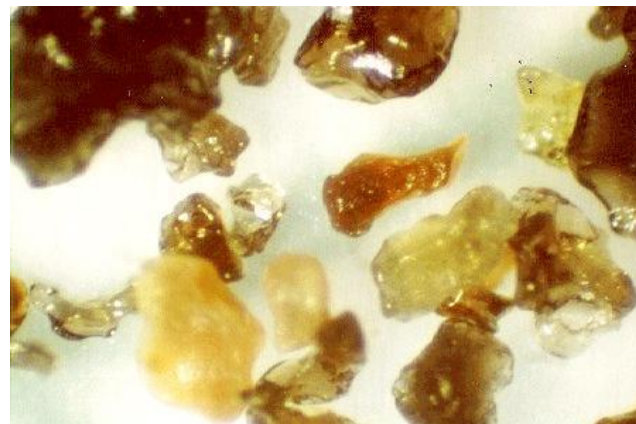
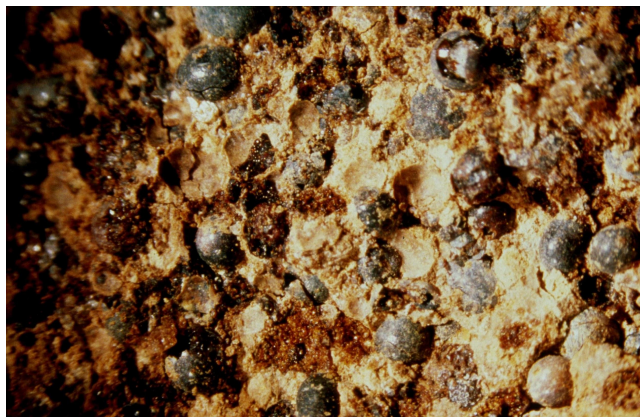
Cicconi et al. (2015) *Amer. Mineral*, 100, 1610-1619

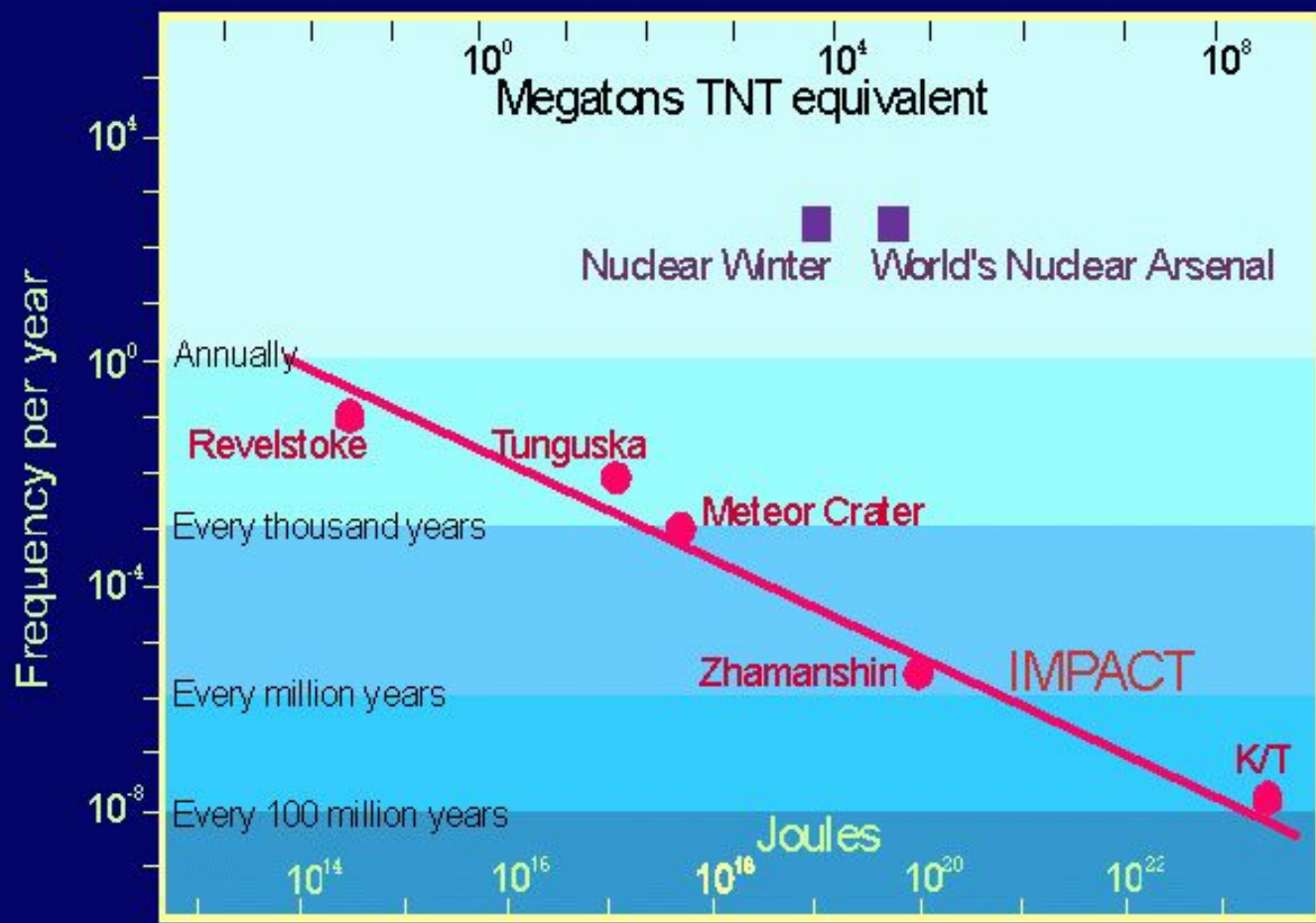
Chixclub (K/Pg), 65My, 270 km ca.



and Cenozoic Coastlines" (Smith et al. 1994)

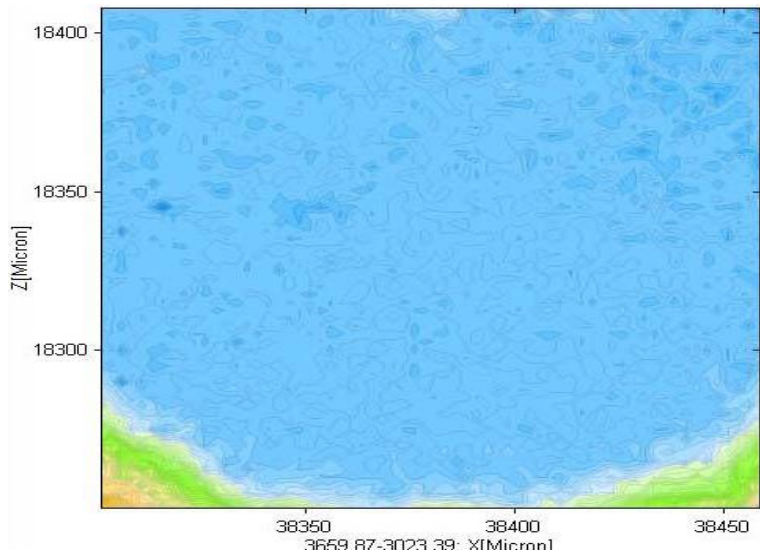






Water content

μ FTIR MAP



Water content in tektites and microtektites

