

# Direct and inverse problems for a class of degenerate parabolic equations arising in climate science

*Piermarco Cannarsa*

Dipartimento di Matematica  
Università di Roma "Tor Vergata"

International Conference  
on **Partial Differential Equations and Applications**  
*in honor of the 70th birthday of Pierangelo Marcati*

Gran Sasso Science Institute, L'Aquila (IT) 19-23 June 2023



Figure: UMI National Congress, Pavia 2019



Figure: Piero's new appointment

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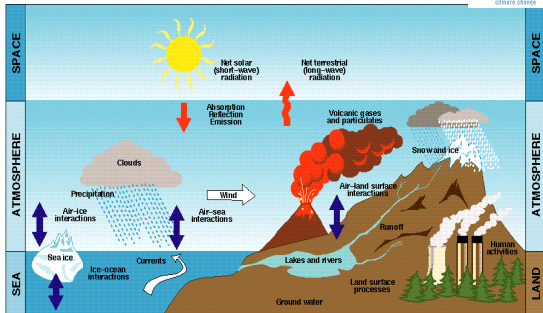
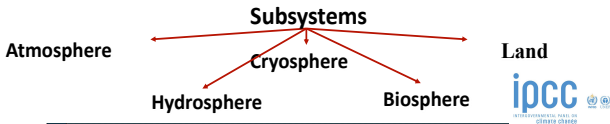
IV. Ongoing research : EBMs with vertical component

# “Climate is what we expect, weather is what we get”

*When we talk about climate, we mean weather averaged over space and time so that local variations and diurnal and random fluctuations have been eliminated*

Mathematics & Climate by H. Kaper and H. Engler

## The Climate System



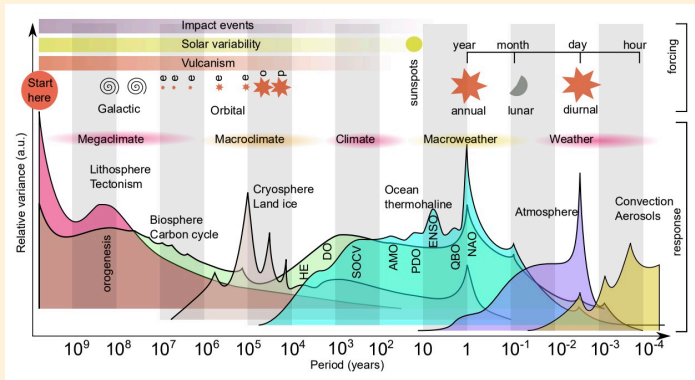
# Climate models

## Purposes

- ▶ Better understanding of past (and future) climate
- ▶ Better understanding the sensitivity to some relevant solar and terrestrial parameters

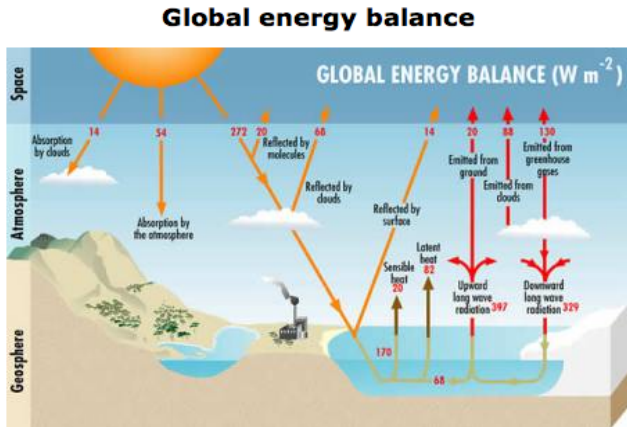
## Main feature

- ▶ Multi-scale nature (in both time and space)



# Energy balance models

## General description

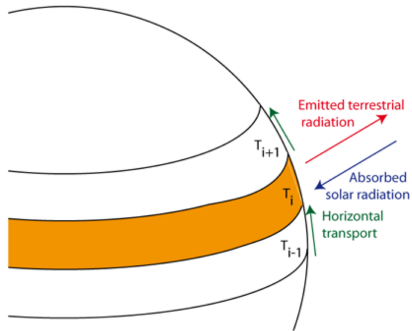


Source: Bureau of Meteorology, [The greenhouse effect and climate change](#), Bureau of Meteorology, 2003, p. 7.

# Energy balance models

introducing diffusion

## Local Energy Budget





## Hierarchy in the class of climate models :

- ▶  $0 - D$  :  $u(t)$ , mean Earth temperature
- ▶  $1 - D$  :  $u(t, \varphi)$  : mean temperature on the latitude circles around the Earth
- ▶  $2 - D$  :  $u(t, m)$  : temperature on Earth surface ( $m \in S^2$ )
- ▶  $3 - D$  : General Circulation Model  $u(t, m, h)$  (coupled with Glaceology, Celestial Mechanics, Geophysics...)

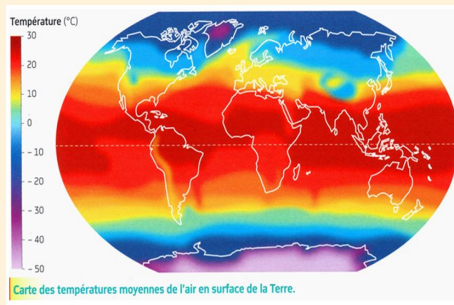


Figure:  $\varphi \in \left(-\frac{\pi}{2}, \frac{\pi}{2}\right)$  parametrizes latitude

# The Budyko (1969) and Sellers (1969) models

Temperature average on the Earth  $u$  satisfies

variation of  $u = +\text{absorbed energy} - \text{reflected energy} + \text{diffusion}$

hence a **reaction-diffusion equation** of the form

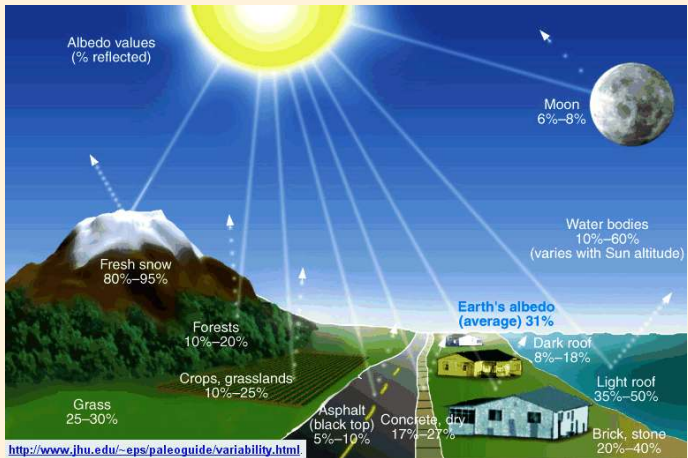
$$c(t, x)u_t - \text{diffusion} = R_a - R_e$$

where

- ▶  $c(t, x)$  : heat capacity,
- ▶  $R_a =$  absorbed solar radiation,  $= QS(t, x)\beta(u)$  :
  - $Q$  : Solar constant,
  - $S(t, x)$  : distribution of solar radiation,
  - $\beta(u)$  : "planetary coalbedo" (= the fraction absorbed according the average temperature),
- ▶  $R_e =$  : emitted radiation (depends on the amount of greenhouse gases, clouds and water vapor in the atmosphere, increases with  $u$ )

# EBMs : absorbed solar radiation

## Albedo



# EBMs : absorbed solar radiation

## Co-albedo and insolation function

$$R_a(t, x) = \beta(u(t, x)) Q(t, x) :$$

- ▶  $Q$  : high-frequency **solar radiation**
- ▶  $\beta(u) = 1 - \alpha(u)$  : **co-albedo** (fraction of absorbed energy) will be assumed to be **piecewise linear**

$$\beta(u) = \begin{cases} \beta_- & \text{for all } u \leq T_- \\ \beta_- + (\beta_+ - \beta_-) \frac{u - T_-}{T_+ - T_-} & \text{for all } u \in [T_-, T_+] \\ \beta_+ & \text{for all } u \geq T_+ \end{cases}$$

for suitable constants  $T_+ > T_- > 0$  and  $\beta_+ > \beta_- > 0$

- ▶  $\alpha$  : **albedo** (fraction of reflected energy);
  - $\alpha$  nonincreasing, from  $\alpha_+$  to  $\alpha_-$  (ice reflects more than non-iced surfaces),
  - **Sellers** :  $\alpha(u)$  smooth / **Budyko** :  $\alpha(u)$  discontinuous,
  - **Bhattacharya-Ghil-Vulis (1982)** :  $\alpha(u, \text{memory effect})$  ;  
(memory effect : interesting to take into account the long response times of the ice sheets to temperature changes).

# EBCMs : diffusion and emitted radiation

► **diffusion** =  $\text{div}(k(\dots)\nabla u)$  :

- $k = k_0$  positive constant, and averaging along the parallels,  $x = \sin(\text{latitude})$  : 1D model, degenerate parabolic equation (and possibly quasilinear) :

$$k_0((1 - x^2)u_x)_x, \quad x \in (-1, 1);$$

- Sellers (1969), Ghil (1976) : 1D,  $k(u)$ ,
- Stone (1972) :  $k(x, \nabla u) = k_1(x)|\nabla u|$  (manifold, rotating atmosphere),
- Diaz (1993) :  $k(x, \nabla u) = k_1(x)|\nabla u|^{p-2}$  (manifold) ;

► **emitted radiation** :

$$\text{Sellers} : R_e = \sigma u^4 \quad / \quad \text{Budyko} : R_e = a + bu,$$

(where  $\sigma$  : Stefan-Boltzmann constant ).

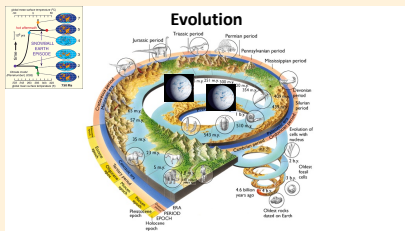
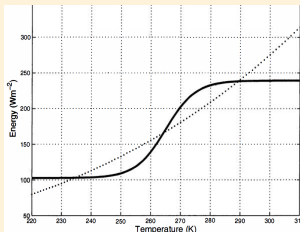
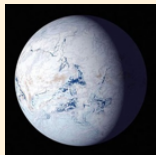
# EBMs : multistability of climate

Snow-ball and warm state

Even the simple ODE

$$u'(t) = \beta(u(t)) - \sigma u^4$$

has been used to prove the multistability of climate



# EBMs : an overview

## Main features

- ▶  $1 - D$  : **degenerate** diffusion coefficient
- ▶  $2 - D$  : on a **manifold**
- ▶ with **nonlinear** source terms, and possibly quasilinear
- ▶ possibly with **discontinuous** coefficients (Budyko)
- ▶ possibly with non local terms (**memory**)

## What has been studied :

- ▶ existence and stability of **multiple** steady states (**Ghil (1976)**)
- ▶ **existence** of solutions, **uniqueness/non uniqueness** (in prescribed classes) (**Diaz (1993), Hetzer (1996, 2011)**)
- ▶ dynamics, **long-time behavior** (**Hetzer (1991)**)
- ▶ **free boundary** value problem : snow lines (**Diaz (1993)**)
- ▶ coeffs : uniqueness, **inverse problems** (**Ghil et al (2014)**)

## II. Sellers model with memory : well-posedness



# 1D Sellers climate model with memory

$$\begin{cases} u_t - (\rho_0(1-x^2)u_x)_x = r(t)q(x)\beta(u) - \varepsilon(u)|u|^3u + \mathbf{f}(\mathbf{H}), \\ \rho_0(1-x^2)u_x = 0, & x = \pm 1, \\ u(s, x) = u_0(s, x), & s \in [-\tau, 0], \end{cases}$$

where

- ▶ 1-D parametrization  $x = \sin(\varphi) \in (-1, 1)$  with  $\varphi =$  the latitude
- ▶ absorbed energy
- ▶ emitted energy
- ▶ **memory term** : to account for long response times of ice sheets to temperature changes (Ghil et al 1982, 2014)

$$\mathbf{H}(\mathbf{t}, \mathbf{x}, \mathbf{u}) = \int_{-\tau}^0 k(s, x)u(t + s, x) ds.$$

# An inverse problem for the Sellers model

- ▶ **Goal** : study an inverse problem that consists in recovering the insulation function  $q(x)$  in the **Sellers model with memory** using partial measurements of the solution,
- ▶ **Difficulties** : degeneracy + nonlinearity + nonlocal,
- ▶ **Results** :
  - well-posedness,
  - uniqueness result under pointwise measurements,
  - Lipschitz stability under localized measurements.

## Sellers model : precise assumptions

- ▶  $\rho(x) = \rho_0(1 - x^2)$ ,  $\rho_0 > 0$ ,  $x \in (-1, 1)$ ,
- ▶  $\beta \in \mathcal{C}^2(\mathbb{R})$ ,  $\beta, \beta', \beta'' \in L^\infty(\mathbb{R})$ ,  $\beta(\cdot) \geq \beta_1 > 0$ ,
- ▶  $q \in L^\infty(I)$ ,
- ▶  $r \in \mathcal{C}^1(\mathbb{R})$ ,  $r, r' \in L^\infty(\mathbb{R})$ ,  $r(\cdot) \geq r_1 > 0$ ,
- ▶  $\varepsilon \in \mathcal{C}^2$ ,  $\varepsilon, \varepsilon', \varepsilon'' \in L^\infty(\mathbb{R})$ ,  $\varepsilon(\cdot) \geq \varepsilon_1 > 0$ ,
- ▶ memory term :
  - kernel  $k \in \mathcal{C}^1([- \tau, 0] \times [-1, 1], \mathbb{R})$ ,
  - nonlinearity  $f \in \mathcal{C}^2(\mathbb{R})$ ,  $f, f', f'' \in L^\infty(\mathbb{R})$ .

# 1D Sellers model : functional setting

- ▶ energy space :  $I = (-1, 1)$

$$V = \left\{ w \in L^2(I) : w \in AC_{loc}(I), \sqrt{\rho}w_x \in L^2(I) \right\} \subset L^p(I), \forall p \geq 1$$

- ▶ operator  $A : D(A) \subset L^2(I) \rightarrow L^2(I)$  in the following way :

$$\begin{cases} D(A) := \{u \in V : \rho u_x \in H^1(I)\} \\ Au := (\rho_0(1-x^2)u_x)_x, \quad u \in D(A) : \end{cases}$$

$(A, D(A))$  is a self-adjoint operator and it is the infinitesimal generator of an analytic and compact semigroup  $\{e^{tA}\}_{t \geq 0}$  in  $L^2(I)$  that satisfies

$$\| \| e^{tA} \| \|_{\mathcal{L}(L^2(I))} \leq 1.$$

(Campiti-Metafune-Pallara (1998))

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## 1D Sellers model : definition of mild solution

$$\begin{cases} \dot{u}(t) = Au(t) + G(t, u) + \mathbf{F}(u^{(t)}) & t \in [0, T] \\ u(s) = u_0(s) & s \in [-\tau, 0], \end{cases} \quad (1)$$

with

$G(t, u)$  = local source terms,  $\mathbf{F}(u^{(t)})$  = memory term

### Definition

Given  $u_0 \in C([-\tau, 0]; V)$ , a function

$$u \in H^1(0, T; L^2(I)) \cap L^2(0, T; D(A)) \cap C([-\tau, T]; V)$$

is called a mild solution of (1) on  $[0, T]$  if  $u(s) = u_0(s)$  for all  $s \in [-\tau, 0]$ , and if for all  $t \in [0, T]$ , we have

$$u(t) = e^{tA} u_0(0) + \int_0^t e^{(t-s)A} (G(s, u) + \mathbf{F}(u^{(s)})) ds$$

## Memory Sellers model : well-posedness result

$$\begin{cases} \dot{u}(t) = Au(t) + G(t,u) + \mathbf{F}(u^{(t)}) & t \in [0, T] \\ u(s) = u_0(s) & s \in [-\tau, 0], \end{cases} \quad (1)$$

### Theorem (C-Malfitana-Martinez (2018))

Let  $u_0$  be such that

$$u_0 \in C([-\tau, 0]; V) \quad \text{and} \quad u_0(0) \in D(A) \cap L^\infty(I).$$

Then, for all  $T > 0$ , (1) has a unique mild solution on  $[0, T]$ .

Proof :

- ▶ local existence (fixed point, contraction)
- ▶ uniqueness (Gronwall's lemma),
- ▶ global existence



# Memory Sellers model : local existence

Functional setting :

- ▶ the space of functions

$$\mathbb{X}_R := \left\{ v \in C([- \tau, t^*]; V) \mid \begin{cases} \|v(t)\|_V \leq R \quad \forall t \in [- \tau, t^*], \\ v(t) = u_0(t) \quad \forall t \in [- \tau, 0], \end{cases} \right\},$$

- ▶ and the associated application

$$\Gamma : \mathbb{X}_R \subset C([- \tau, t^*]; V) \rightarrow C([- \tau, t^*]; V)$$

defined by  $\Gamma(u)(t) := u_0(t)$  for  $t \in [- \tau, 0]$ , and

$$\Gamma(u)(t) := e^{t\tilde{A}}u_0(0) + \int_0^t e^{(t-s)\tilde{A}} \left[ \tilde{G}(s, u(s)) + F(u(s)) \right] ds$$

for  $t \in [0, t^*]$  (and  $\tilde{A} = A - Id$  strictly dissipative) :

$\Gamma$  well-defined,  $\Gamma(\mathbb{X}_R) \subset \mathbb{X}_R$  and  $\Gamma$  contraction of  $t^* > 0$  small.

Essential tool (Pazy) :

$$\|(-\tilde{A})^{1/2} e^{t\tilde{A}}\|_{\mathcal{L}(L^2(I))} \leq \frac{c}{\sqrt{t}}.$$

## Memory Sellers model : uniqueness

$u, \tilde{u}$  solutions on  $[0, T_0]$  : the difference  $w = u - \tilde{u}$  solves

$$\begin{cases} w_t - (\rho w_x)_x = G(t, u) - G(t, \tilde{u}) + f(H) - f(\tilde{H}), \\ \rho w_x = 0, \quad t \in (0, T_0), x = \pm 1, \\ w(s, x) = 0, \quad s \in [-\tau, 0], x \in (-1, 1), \end{cases}$$

and then

$$W(T') := \int_0^{T'} \|w(T)\|_{L^2(-1,1)}^2 dT$$

is nondecreasing and satisfies (integration by parts, estimates)

$$W(T') \leq C \int_0^{T'} W(T) dT.$$

$$\text{Gronwall} \quad \implies \quad W = 0 \quad \implies \quad w = 0 \quad \implies \quad u = \tilde{u}.$$

## Memory Sellers model : global existence

Maximal existence time :

$$T^*(u_0) := \sup\{T \geq 0 \text{ s.t. (1) has a mild solution on } [0, T]\},$$

Then

$$u_0 \in C([- \tau, 0]; V), u_0(0) \in D(A) \cap L^\infty(I) \implies T^*(u_0) = +\infty :$$

based on the following boundedness property :

### Theorem

Consider  $u_0 \in C([- \tau, 0]; V)$  and  $u_0(0) \in D(A) \cap L^\infty(I)$ ,  $T > 0$  and  $u$  a mild solution of (1) defined on  $[0, T]$ . Let us denote

$$M_1 := \left( \frac{\|q\|_{L^\infty(I)} \|r\|_{L^\infty(\mathbb{R})} \|\beta\|_{L^\infty(\mathbb{R})} + \|f\|_{L^\infty(\mathbb{R})}}{\varepsilon_1} \right)^{\frac{1}{4}}$$

and  $M := \max\{\|u_0(0)\|_{L^\infty(I)}, M_1\}$ . Then  $u$  satisfies

$$\|u\|_{L^\infty((0, T) \times I)} \leq M.$$

### III. Sellers model : reconstruction of insolation function

# 1D Memory Sellers model : uniqueness/stability of the insolation function ?

$$\begin{cases} u_t - (\rho(x)u_x)_x = r(t)q(x)\beta(u) - \varepsilon(u)|u|^3u + f(H) \\ \rho(x)u_x = 0, \quad x \in \partial I, \\ u(s, x) = u_0(s, x), \quad s \in [-\tau, 0], x \in I \end{cases} \quad (S)$$

$$\begin{cases} \tilde{u}_t - (\rho(x)\tilde{u}_x)_x = r(t)\tilde{q}(x)\beta(\tilde{u}) - \varepsilon(\tilde{u})|\tilde{u}|^3\tilde{u} + f(\tilde{H}) \\ \rho(x)\tilde{u}_x = 0, \quad x \in \partial I \\ \tilde{u}(s, x) = \tilde{u}_0(s, x), \quad s \in [-\tau, 0], x \in I \end{cases} \quad (\tilde{S})$$

$u = \tilde{u}$  on a "small" set  $\implies q = \tilde{q}$  ?

$u - \tilde{u}$  "small on a small set"  $\implies q - \tilde{q}$  small ?

# Motivation for inverse problems (Roques et al 2014)

- ▶ Goal of the Energy Balance Models (with Memory) : toy models to understand the evolution of climate
- ▶ With suitable tuning of the parameters : EBMs simulations give reasonable results for the observed present climate (North-Mengel-Short (1983))
- ▶ Once fitted, EBM(M) can be used to estimate the temporal response patterns to various scenarios (climate change).
- ▶ BUT in practice, parameters cannot be measured directly (intertwined effects of several physical processes). So, one takes measurements of the solution and uses such measurements to reconstruct parameters (Yamamoto-Zou (2001), Roques et al (2014)).

# Inverse source problem for the linear heat equation

## Various approaches and results

- ▶ **"Simple" models** (constant coefficients/depending only on  $x$  or only on  $t...$ ) : elegant and sharp techniques :
  - Fourier series (moment method, biorthogonal families),
  - Laplace transform,
  - Volterra integral equations...

↪ sharp results (explicit formula of the solution...)

Cannon 1968, Lorenzi-Sinestrari (1988), Lorenzi (1989...), Bukhgeim (1993), Gentili (1991), Grasselli (1992), Yamamoto (1993), Janno-Wolfersdorf (1996), Choulli-Yamamoto (2006)...
- ▶ **nonlinear models** (or coefficients in  $x$  and  $t$ ) : local/global Carleman estimates ↪ uniqueness, Holder/Lipschitz stability : Bukhgeim/Klibanov 1981, Klibanov (1992), Isakov (1990, 1998...), Imanuvilov/Yamamoto (1998)
- ▶ **Use of analyticity properties** ↪ uniqueness under measurements at one point (in  $1 - D$ ) (Roques-Cristofol (2010), Roques-Checkroun-Cristofol-Soubeyrand-Ghil (2014))

# Inverse source problem for evolution equations

## Literature on the subject

### **Founding papers using GCE :**

Puel/Yamamoto 1996 + 1997 (linear wave equation)

Imanuvilov/Yamamoto 1998 (linear heat equation)

### **More on Lipschitz stability for parabolic equations :**

Yamamoto/Zou 2001 (simultaneous reconstruction of 2 quantities)

Cristofol/Gaitan/Ramoul 2006 (systems)

Benabdallah/Dermenjian/Le Rousseau 2007 + Benabdallah/Gaitan/Le

Rousseau 2009 (discontinuous diffusion coefficient)

C/Tort/Yamamoto 2010 (degenerate diffusion coefficient)

Ignat/Pazoto/Rosier 2012 (networks)

### **Lipschitz stability for other types of equations :**

Hyperbolic equations : Imanuvilov/Yamamoto 2001,  
Komornik/Yamamoto 2002, Bellassoued/Yamamoto 2006, Liu/Triggiani  
2011

Schrodinger equation : Baudouin/Puel 2002, Mercado/Osses/Rosier  
2008, Liu/Triggiani 2011 ...



# Uniqueness under pointwise measurements

## Assumptions

(H1) **admissible initial conditions**

$$\mathcal{U}^{(pt)} = C^{1,2}([-\tau, 0] \times [-1, 1])$$

(H2) **admissible coefficients**

$$\mathcal{Q}^{(pt)} := \{q \text{ is Lipschitz-continuous and piecewise analytic on } I\}$$

(H3) **admissible memory kernels**

$$\exists \delta \in (0, \tau) \quad \text{such that} \quad k(s, \cdot) \equiv 0 \quad \forall s \in [-\delta, 0]$$

# Uniqueness under pointwise measurements

## Main result

### Theorem (C-Malfitana-Martinez)

Consider

- ▶ two insulation functions  $q, \tilde{q} \in \mathcal{Q}^{(pt)}$
- ▶ an initial condition  $u_0 = \tilde{u}_0 \in \mathcal{U}^{(pt)}$

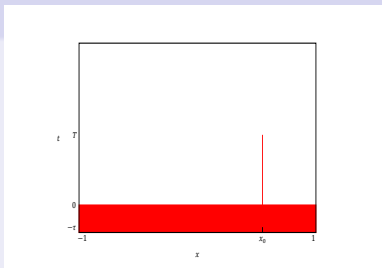
Assume that

- ▶ the memory kernel is admissible
- ▶  $r$  and  $\beta$  are positive

If  $u$  and  $\tilde{u}$  are the solutions of  $(S)$  and  $(\tilde{S})$ , respectively, and there exists  $x_0 \in I$  and  $T > 0$  such that

$$\forall t \in (0, T), \quad u(t, x_0) = \tilde{u}(t, x_0), \text{ and } u_x(t, x_0) = \tilde{u}_x(t, x_0),$$

then  $q \equiv \tilde{q}$  on  $(-1, 1)$



This extends a result by [Roques-Checkroun-Cristofol-Soubeyrand-Ghil \(2014\)](#)

# 1D Memory Sellers model : Lipschitz stability

Set-up

$$\begin{cases} u_t - (\rho(x)u_x)_x = r(t)q(x)\beta(u) - \varepsilon(u)|u|^3u + f(H) \\ \rho(x)u_x = 0, \quad x \in \partial I, \\ u(s, x) = u_0(s, x), \quad s \in [-\tau, 0], x \in I \end{cases} \quad (S)$$

$$\begin{cases} \tilde{u}_t - (\rho(x)\tilde{u}_x)_x = r(t)\tilde{q}(x)\beta(\tilde{u}) - \varepsilon(\tilde{u})|\tilde{u}|^3\tilde{u} + f(\tilde{H}) \\ \rho(x)\tilde{u}_x = 0, \quad x \in \partial I, \\ \tilde{u}(s, x) = \tilde{u}_0(s, x), \quad s \in [-\tau, 0], x \in I \end{cases} \quad (\tilde{S})$$

**Goal** : to prove that

$$\|q - \tilde{q}\| \leq C \|u - \tilde{u}\|$$

# 1D Memory Sellers model : Lipschitz stability

## Assumptions

- ▶ **admissible initial conditions** : given  $M > 0$ ,

$$\mathcal{U}_M^{(loc)} = \left\{ u_0 \in C([- \tau, 0]; V \cap L^\infty(-1, 1)) : u_0(0) \in D(A), Au_0(0) \in L^\infty(I), \right. \\ \left. \sup_{t \in [- \tau, 0]} (\|u_0(t)\|_V + \|u_0(t)\|_{L^\infty}) + \|Au_0(0)\|_{L^\infty(I)} \leq M \right\}$$

- ▶ **admissible coefficients** : given  $M' > 0$ ,

$$\mathcal{Q}_{M'}^{(loc)} := \{q \in L^\infty(I) : \|q\|_{L^\infty(I)} \leq M'\}$$

- ▶ **admissible memory kernels** : the same **support condition**

$$\exists \delta \in (0, \tau) \quad \text{such that} \quad k(s, \cdot) \equiv 0 \quad \forall s \in [-\delta, 0]$$

# 1D Memory Sellers model : Lipschitz stability

## Main result

### Theorem (C–Malfitana–Martinez (2018))

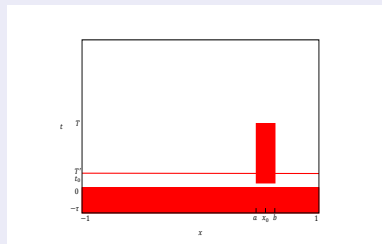
Let  $0 \leq t_0 < T' < T$ , let  $M, M' > 0$ ,  
and suppose

$$0 < T' < \delta$$

Then there exists

$C(t_0, T', T, M, M') > 0$  such that, for  
all  $u_0, \tilde{u}_0 \in \mathcal{U}_M^{(loc)}$  and all  $q, \tilde{q} \in \mathcal{Q}_{M'}^{(loc)}$ ,  
the solutions  $u$  and  $\tilde{u}$  of (S) and ( $\tilde{S}$ ),  
respectively, satisfy

$$\begin{aligned} \|q - \tilde{q}\|_{L^2(I)}^2 &\leq C \left( \|u(T') - \tilde{u}(T')\|_{D(A)}^2 \right. \\ &\quad \left. + \|u_t - \tilde{u}_t\|_{L^2((t_0, T) \times (a, b))}^2 + \|u_0 - \tilde{u}_0\|_{C([-\tau, 0], V)}^2 \right) \end{aligned}$$



Extension of Tort–Vancostenoble (2012)

# 1D Memory Sellers model : Lipschitz stability

## Main tools for proof

- ▶ to reduce the problem to some (non standard) inverse source problem for a linear equation
- ▶ to follow the method by Imanuvilov-Yamamoto (1998) for inverse source problems
- ▶ to use adapted Global Carleman Estimates for degenerate parabolic equations C – Martinez-Vancostenoble (2008)
- ▶ appeal to maximum principles to deal with nonlinear terms

# 1D Memory Sellers model : Lipschitz stability

Steps 1 and 2 of the proof

Step 1 to identify the problem satisfied by the difference  $w = u - \tilde{u}$  as

$$\begin{cases} w_t - (\rho w_x)_x = K^* + K + \tilde{K} + K^h, & t > 0, x \in (-1, 1), \\ \rho w_x = 0, & x = \pm 1, \\ w(s, x) = u_0(s, x) - \tilde{u}_0(s, x), & s \in [-\tau, 0], x \in (-1, 1), \end{cases}$$

where

$$\begin{aligned} K^*(t, x) &= r(t)(q(x) - \tilde{q}(x))\beta(u), & K(t, x) &= r(t)\tilde{q}(x)(\beta(u) - \beta(\tilde{u})), \\ \tilde{K} &= -\varepsilon(u)|u|^3u + \varepsilon(\tilde{u})|\tilde{u}|^3\tilde{u}, & K^h &= f(H) - f(\tilde{H}) \end{aligned}$$

Step 2  $K^*$  satisfies (Imanuvilov-Yamamoto (1998))

$$\exists C_0 > 0 \text{ s.t. } \forall t, x, \quad \left| \frac{\partial K^*}{\partial t}(t, x) \right| \leq C_0 |K^*(T', x)|$$

(this follows from the fact that  $u_t \in L^\infty((0, T) \times I)$ )

# 1D Memory Sellers model : Lipschitz stability

## Carleman estimate and conclusion

**Step 3 Carleman estimate for  $z := w_t$**  : there exists a smooth weight function  $\theta : (t_0, T) \rightarrow (0, \infty)$ , with  $\theta(t) \rightarrow \infty$  as  $t \downarrow t_0$  and  $T \uparrow T$ , such that

$$\begin{aligned} & \int_{t_0}^T \int_{-1}^1 \left( R^3 \theta^3 (1-x^2) z^2 + R \theta (1-x^2) z_x^2 + \frac{1}{R \theta} z_t^2 \right) e^{-2R\sigma} \\ & \leq C \int_{t_0}^T \int_{-1}^1 (K_t^*)^2 e^{-2R\sigma} + C \int_{t_0}^T \int_a^b R^3 \theta^3 z^2 e^{-2R\sigma} \\ & \quad + C \|w(T')\|_{L^2(I)}^2 + C \|u_0 - \tilde{u}_0\|_{C([- \tau, 0], V)}^2 \end{aligned}$$

**Step 4** to estimate  $\|q - \tilde{q}\|$  we use an upper bound for

$$\begin{aligned} K^*(T') &= r(T')(q(x) - \tilde{q}(x))\beta(u) \\ &= z(T') - (\rho w_x)_x(T') - K(T') - \tilde{K}(T') - K^h(T') \end{aligned}$$

that is

$$\begin{aligned} & \int_{-1}^1 |K^*(T')|^2 e^{-2R\sigma(T')} dx \leq C \|w(T')\|_{D(A)}^2 + \\ & + C \int_{-1}^1 \left( \underbrace{|z(T')|^2}_{(CE)} + \underbrace{|K(T')|^2 + |\tilde{K}(T')|^2}_{C \|w(T')\|_{L^2(I)}^2} + \underbrace{|K^h(T')|^2}_{C \|u_0 - \tilde{u}_0\|_{C([- \tau, 0], V)}^2} \right) e^{-2R\sigma(T')} dx \end{aligned}$$



## IV. Ongoing research : EBMs with vertical component

# EBMs with vertical component

joint work with V. Lucarini, P. Martinez, C. Urbani, J. Vancostenoble

$T_a$  : temperature of an **atmosphere layer**

$T_s$  : **surface** temperature of the Earth

$$\begin{cases} \gamma_a \left[ \frac{\partial T_a}{\partial t} - k_a \frac{\partial}{\partial x} \left( (1-x^2) \frac{\partial T_a}{\partial x} \right) \right] = -\lambda(T_a - T_s) + \varepsilon_a \sigma_B |T_s|^3 T_s - 2\varepsilon_a \sigma_B |T_a|^3 T_a + \mathcal{R}_a \\ \gamma_s \left[ \frac{\partial T_s}{\partial t} - k_s \frac{\partial}{\partial x} \left( (1-x^2) \frac{\partial T_s}{\partial x} \right) \right] = -\lambda(T_s - T_a) - \sigma_B |T_s|^3 T_s + \varepsilon_a \sigma_B |T_a|^3 T_a + \mathcal{R}_s \\ (1-x^2) \frac{\partial T_a}{\partial x} \Big|_{x=\pm 1} = 0 = (1-x^2) \frac{\partial T_s}{\partial x} \Big|_{x=\pm 1} \quad T_a(0, x) = T_a^{(0)}(x), \quad T_s(0, x) = T_s^{(0)}(x) \end{cases}$$

**Exchange of energy between the layers :**

- ▶ linear term : non-radiative vertical exchanges of energy due to the action of the geophysical fluids
- ▶ nonlinear term : emission of infrared radiation by one level being captured by the other layer

**Relevant constants :**

- ▶  $\lambda \geq 0$  : coupling parameter for vertical exchanges
- ▶  $\varepsilon_a \in [0, 1]$  : absorptivity (depends on greenhouse gases  $CO_2$ ,  $CH_4$ )

# Well-posedness of the ODE problem

$$(2LEBM) \quad \begin{cases} \gamma_a \frac{\partial T_a}{\partial t} = -\lambda(T_a - T_s) + \varepsilon_a \sigma_B |T_s|^3 T_s - 2\varepsilon_a \sigma_B |T_a|^3 T_a + \mathcal{R}_a, \\ \gamma_s \frac{\partial T_s}{\partial t} = -\lambda(T_s - T_a) - \sigma_B |T_s|^3 T_s + \varepsilon_a \sigma_B |T_a|^3 T_a + \mathcal{R}_s, \\ T_a(0) = T_a^{(0)}, \quad T_s(0) = T_s^{(0)} \end{cases}$$

**Assumptions :**

- ▶  $\lambda \geq 0$ ,  $q > 0$ ,  $\sigma_B > 0$ ,  $\varepsilon_a \in (0, 2)$
- ▶  $\beta_a, \beta_s : \mathbb{R} \rightarrow \mathbb{R}$  globally Lipschitz,  $\beta_a \geq 0$  and  $\beta_s > 0$ , and

$$\mathcal{R}_a = q\beta_a(T_a), \quad \mathcal{R}_s = q\beta_s(T_s).$$

- ▶  $T_a^{(0)} \geq 0$ ,  $T_s^{(0)} \geq 0$

## Proposition

(2LEMB) has unique solution, defined and bounded for any  $t \in [0, +\infty)$ .

Moreover

$$\forall t \in (0, +\infty), \quad T_a(t) > 0 \quad \text{and} \quad T_s(t) > 0.$$

# Asymptotic behaviour of solutions

## Definition

A  $C^1$  system of ODE on  $\mathbb{R}^n$   $\frac{d}{dt}x_i = F_i(x_1, \dots, x_n) = F_i(x)$ ,  $i = 1, \dots, n$  is *competitive* if

$$\frac{\partial}{\partial x_j} F_i(x) \leq 0 \quad i \neq j$$

and **cooperative** if the reverse inequalities hold.

For  $\varepsilon_a \in (0, 2)$  and  $\lambda \geq 0$

$$T_s \mapsto \frac{1}{\gamma_a} \left[ \lambda T_s + \varepsilon_a \sigma_B |T_s|^3 T_s - \lambda T_a - 2\varepsilon_a \sigma_B |T_a|^3 T_a + \mathcal{R}_a(T_a) \right]$$

$$T_a \mapsto \frac{1}{\gamma_s} \left[ \lambda T_a - \sigma_B |T_s|^3 T_s - \lambda T_s + \varepsilon_a \sigma_B |T_a|^3 T_a + \mathcal{R}_s(T_s) \right]$$

are nondecreasing on  $[0, +\infty)$

$\implies$  our system is cooperative

$\implies$  [Smith] any initial condition  $(T_a^{(0)}, T_s^{(0)})$  converges to an equilibrium.

## Case $\lambda = 0$ and $\mathcal{R}_a = 0$

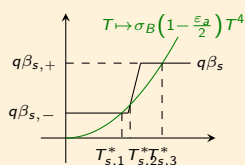
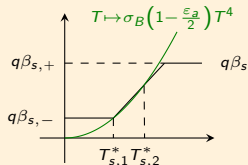
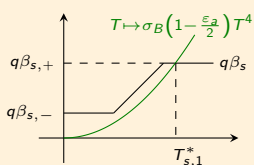
Equilibrium points are solutions of

$$\begin{cases} \varepsilon_a \sigma_B T_s^4 - 2\varepsilon_a \sigma_B T_a^4 = 0, \\ -\sigma_B T_s^4 + \varepsilon_a \sigma_B T_a^4 + \mathcal{R}_s(T_s) = 0 \end{cases}$$

that is equivalent to solve

$$\sigma_B \left(1 - \frac{\varepsilon_a}{2}\right) T_s^4 = q\beta_s(T_s)$$

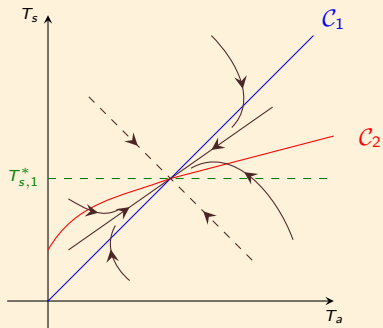
$\Rightarrow$  one, two or three possible equilibria



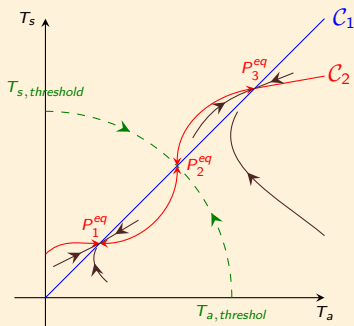
depending on parameters  $\sigma_B$ ,  $\varepsilon_a$ ,  $q$ .

# Nature of the equilibrium points ( $\lambda = 0, \mathcal{R}_a = 0$ )

One equilibrium :  $(T_{a,1}^*, T_{s,1}^*)$   
**asymptotically stable**



Three equilibria :  $(T_{a,1}^*, T_{s,1}^*)$ ,  
 $(T_{a,3}^*, T_{s,3}^*)$  **asymptotically**  
**exponentially stable**,  $(T_{a,2}^*, T_{s,2}^*)$   
**unstable**



## Case $\lambda > 0$ and $\mathcal{R}_a = 0$

Equilibrium points are solutions of

$$\begin{cases} -\lambda(T_a - T_s) + \varepsilon_a \sigma_B |T_s|^3 T_s - 2\varepsilon_a \sigma_B |T_a|^3 T_a = 0, \\ -\lambda(T_s - T_a) - \sigma_B |T_s|^3 T_s + \varepsilon_a \sigma_B |T_a|^3 T_a + q\beta(T_s) = 0. \end{cases}$$

### Lemma

- ▶ *There exists at least one warm and one cold equilibrium.*
- ▶ *There exists at most five equilibria.*
- ▶ *Fixed  $\lambda > 0$  and  $\varepsilon_a \in (0, \varepsilon_{a,0})$  (with  $\varepsilon_{a,0} > 1.99$ ), there exist at most three equilibrium points.*
- ▶ *Any warm (resp. a cold) equilibrium is asymptotically exponentially stable.*

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# Dependence of equilibria on $\varepsilon_a$

## Proposition

Fixed  $\lambda \geq 0$ , let  $(T_a^{\text{eq}, \varepsilon_a^*}, T_s^{\text{eq}, \varepsilon_a^*})$  be an asymptotically exponentially stable warm [resp. cold] equilibrium point with  $\varepsilon_a = \varepsilon_a^*$ .

Then, there exists a unique asymptotically exponentially stable warm [resp. cold] equilibrium  $(T_a^{\text{eq}, \varepsilon_a}, T_s^{\text{eq}, \varepsilon_a})$  for  $\varepsilon_a$  close to  $\varepsilon_a^*$  and the following monotonicity property holds :

- ▶  $\varepsilon_a \mapsto T_s^{\text{eq}, \varepsilon_a}$  is increasing : the surface temperature of the equilibrium increases as  $\varepsilon_a$  increases ;
- ▶  $\varepsilon_a \mapsto T_a^{\text{eq}, \varepsilon_a}$  is increasing if  $\varepsilon_a^* \in (1, 2)$  : the atmosphere temperature of the equilibrium increases as  $\varepsilon_a$  increases.

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Fixed  $\varepsilon_a \in (0, 2)$ , let  $(T_a^{\text{eq}, \lambda^*}, T_s^{\text{eq}, \lambda^*})$  be an asymptotically exponentially stable warm [resp. cold] equilibrium with  $\lambda = \lambda^* \geq 0$ .

Then, there exists a unique asymptotically exponentially stable equilibrium point  $(T_a^{\text{eq}, \lambda}, T_s^{\text{eq}, \lambda})$  for  $\lambda$  close to  $\lambda^*$  and the following monotonicity properties are satisfied :

- ▶  $\lambda \mapsto T_s^{\text{eq}, \lambda}$  is decreasing : the surface temperature of the equilibrium decreases as  $\lambda$  increases ;
- ▶  $\lambda \mapsto T_a^{\text{eq}, \lambda}$  is
  - increasing if  $\varepsilon_a \in (0, 1)$ ,
  - decreasing if  $\varepsilon_a \in (1, 2)$ .

Hence, the atmosphere temperature of an equilibrium is monotone with respect to  $\lambda$ , and monotonicity depends on  $\varepsilon_a$ .

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Hence, the atmosphere temperature of an equilibrium is monotone with respect to  $\lambda$ , and monotonicity depends on  $\varepsilon_a$ .

## Future work directions

- ▶ Inverse problems for parameter reconstruction ( $\varepsilon_a$ ,  $\lambda$ ,  $q$ )
- ▶ Analysis of the PDE system
- ▶ Extension to a variable solar radiation  $Q(t, x) = r(t)q(x)$ , with  $r$  positive and periodic - allowing for seasonal cycle - and  $q$  the latitudinal-dependent insolation function
- ▶ Extension to a space-dependent absorptivity  $\varepsilon_a(x)$





*Happy Birthday Piero!*